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**Acoustic Tomography in Boreholes
using an Algebraic Reconstruction Technique.**

by

Kingsley L. Smith

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A thesis submitted to the Faculty and the Board of Trustees of the Colorado School of Mines in partial fulfillment of the requirements for the degree of Master of Science (Geophysics).

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ABSTRACT

An algebraic reconstruction technique (ART) is described for the seismic tomography of velocities from the travel times of a multiple offset vertical seismic profile. ART concentrates on the production of a reconstructed field whose projected data (travel times) agree with the observed data. This reconstructed field is modified by altering the data for each ray such that when this data is back-projected, the new image agrees with the original data. Because the paths of the rays must be known to calculate the expected travel times, the problem is linearized by using raypath approximations as determined from either a constant or a linear $c(z)$ velocity medium.

Imaging of synthetic data revealed that the orientation of the anomaly affects both the rate of convergence and the resolution of the reconstructed field. Some smoothing of the velocity anomalies occurred along the direction of the rays.

Noisy data sets developed problems in the reconstructed velocity field. Huge single point anomalies appeared along the model's edge in the reconstructed image. Elimination of these errant anomalies was necessary to obtain a reasonable velocity reconstruction. Because isotropy was assumed, the algorithm poorly reconstructed data which was collected over the strongly anisotropic Pierre Shale.

TABLE OF CONTENTS

Abstract iii

Table of Contents iv

List of Figures and Tables v

Notation vii

Acknowledgements ix

1. Introduction 1

2. Algebraic Reconstruction Technique - ART 4

2.1 Constant Background 6

2.2 Constant Background - Iterative Algorithm. 10

2.3 Linear $c(z)$ Medium 12

2.4 Linear $c(z)$ Medium - Iterative Algorithm 19

3. Testing the Algorithms 21

4. Results 22

5. Pierre Shale Experiment 32

6. Conclusions 56

References 58

Appendix A: Ray Parameter Derivation 60

Appendix B: Travel Time Calculation 63

Appendix C: Computer Program Listing 65

LIST OF FIGURES AND TABLES

Figure 1.	Multiple offset VSP geometry.	5
Figure 2.	Modeling the medium using overlapping circles.	8
Figure 3.	Derivation of the local radius of curvature.	13
Figure 4.	Origin determination.	16
Figure 5.	Arc length determination.	18
Figure 6.	Velocity increasing with x.	23
Figure 7.	Velocity increasing with x.	24
Figure 8.	Dipping model.	25
Figure 9.	Model of a low velocity anomaly.	26
Figure 10.	Velocity increasing with x and z.	27
Figure 11.	Two-layered model.	28
Figure 12.	Dipping model.	33
Figure 13.	High velocity anomaly within a linear trend.	34
Figure 14.	Faulted model.	35
Table 1.	Travel times in msec.	36
Figure 15.	Field experiment layout.	37
Figure 16.	Typical waveforms.	38
Figure 17.	Acoustic tomography from the Pierre Shale VSP data. . . .	40
Figure 18.	Acoustic tomography from the Pierre Shale VSP data. . . .	42
Figure 19.	Acoustic tomography from the Pierre Shale VSP data. . . .	45
Figure 20.	Acoustic tomography from the Pierre Shale VSP data. . . .	47
Figure 21.	Acoustic tomography from the Pierre Shale VSP data. . . .	49

Figure 22. Acoustic tomography from the Pierre Shale VSP data. . . . 51
Figure 23. Travel time curves. 54

NOTATION

α	Angle formed by the length of the raypath.
c_0	Velocity at the surface.
$c(z)$	Velocity.
d^{mn}	Sum of chord lengths.
e_{ij}^{mn}	Residual error time for chord length.
E_{ij}	Residual error time for entire ray.
i	Angle of incidence.
ij	Subscript - from the i 'th source to the j 'th receiver.
l_{ij}^{mn}	Chord length.
L_{ij}	Total length of the ray.
mn	Superscript for parameter circle around point m,n .
m_0	Slope of the linear $c(z)$ velocity.
p	Ray parameter.
r	Radius of any parameter circle.
R	Radius of curvature for the ray.
s^{mn}	Slowness matrix.
\tilde{s}^{mn}	Slowness correction matrix.
t_{ij}^{mn}	Residence time.
T_{ij}	Calculated travel time for the ray ij .
\tilde{T}_{ij}	Observed travel time for the ray ij .
w_{ij}^{mn}	Weighting factor.
x_i, z_j	Offset to source, depth to receiver.

x_m, z_n	Offset and depth to parameter point mn.
x_o, z_o	Origin of the radius of curvature for the ray.

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1. INTRODUCTION

One of the many goals of vertical seismic profiling (VSP) is to determine the lateral variations which occur away from the borehole. These lateral changes may reflect porosity, permeability, or facies changes within the rock unit. Such occurrences greatly influence the migration and accumulation of petroleum. Therefore, methods are needed to delineate these discrepancies between the borehole wall and the surrounding medium.

This investigation was geared at determining velocity inhomogeneities using only the first arrival times from a multiple offset VSP. The travel time through any medium is the integral of the slowness (reciprocal of the velocity) along the raypath. Unfortunately, this raypath itself is dependent upon the velocities within the medium. Thus, as posed, the problem is non-linear.

Linearizing this problem requires that the raypaths are known, which, in turn, necessitates a prior knowledge of the velocity structure. However, if we assume some general velocity trend for the medium, we can approximate the paths which the rays take. Then, the problem becomes one of finding the velocity perturbation from the assumed background velocity. Furthermore, we assume that all raypaths are not

influenced by this perturbation velocity; this assumption holds true for small perturbations. Since the rays are independent of the velocity perturbation, the problem becomes linear.

This research utilized two background models:

- 1) Constant velocity medium in which the rays form straight lines from the source to the receiver.
- 2) Linear $c(z)$ velocity medium where the raypaths follow arcs along a circle.

The residual time is defined as the observed time minus the expected time as calculated from the velocity model. Fawcett (1983) utilized a Radon transform method for his seismic tomography of the slowness field from reflected residual travel times. Neumann (1981) used a least-squares inversion of residual travel times in his study on reflection seismics. Christofferson and Husebye (1979) applied a least-squares approach on P-wave residual times in a three-dimensional case. In contrast, McMechan (1983) and Mason (1981) opted for an algebraic reconstruction technique (ART). I have opted for the last approach, as well.

In comparison to a least-squares inversion, ART possesses several advantages:

- 1) ART programs are computationally faster and they are easier to program.
- 2) Constraints are easily incorporated into the program to accommodate any prior knowledge of the medium.
- 3) ART can be easily applied to any source/receiver geometry without difficulties.

ART has its origins within the medical field. Gordon, Bender, and Herman (1970) first introduced ART for the reconstruction of images from x-ray pictures. Later, Herman, Lent, and Rowland (1973) improved the early ART algorithms and Gordon (1974) summarized the various ART developments within the medical profession.

Several authors worked on reconstruction techniques which are applicable to the more general reconstruction problem rather than the specific x-ray case. Mersereau and Oppenheim (1974) and Mueller, Kaveh, and Wade (1979) applied ART to density reconstruction problems using the Fourier domain. Horn (1978) developed a density reconstruction algorithm for any arbitrary ray sampling scheme.

Geophysical tomography is relatively new. Dines and Lytle (1979) reconstructed pictures of electromagnetic properties in the region between a pair of boreholes. By assuming only straight raypaths, seismic velocity structures were obtained by both Mason (1981) and McMechan (1983). Mason concerned himself with inverting seismic data

shot between two boreholes, while McMechan imaged data shot from a borehole to a second borehole as well as to the surface.

The VSP model used in this paper consisted of ten point sources, which were equally spaced at 100 foot intervals along the surface. Similarly, a 100 foot spacing was used for the ten receivers placed down the borehole. No source or receiver resided at the top of the borehole (see Fig. 1). For a coordinate system, whose origin rested at the surface expression of the borehole, the x variable coincided with the surface offset distance and the z variable reflected the depth below the surface.

2. ALGEBRAIC RECONSTRUCTION TECHNIQUE - ART

The ART algorithm used here was adapted from Mason (1981). It involves a nine step process:

- 1) The medium is modeled by casting a set of overlapping circles on a uniform grid.
- 2) The total length of ray segments about a parameter is the sum of all the chord lengths formed by rays which pass through any given parameter circle.
- 3) The travel times associated with each of these chord lengths are determined by the velocity background model and summed.

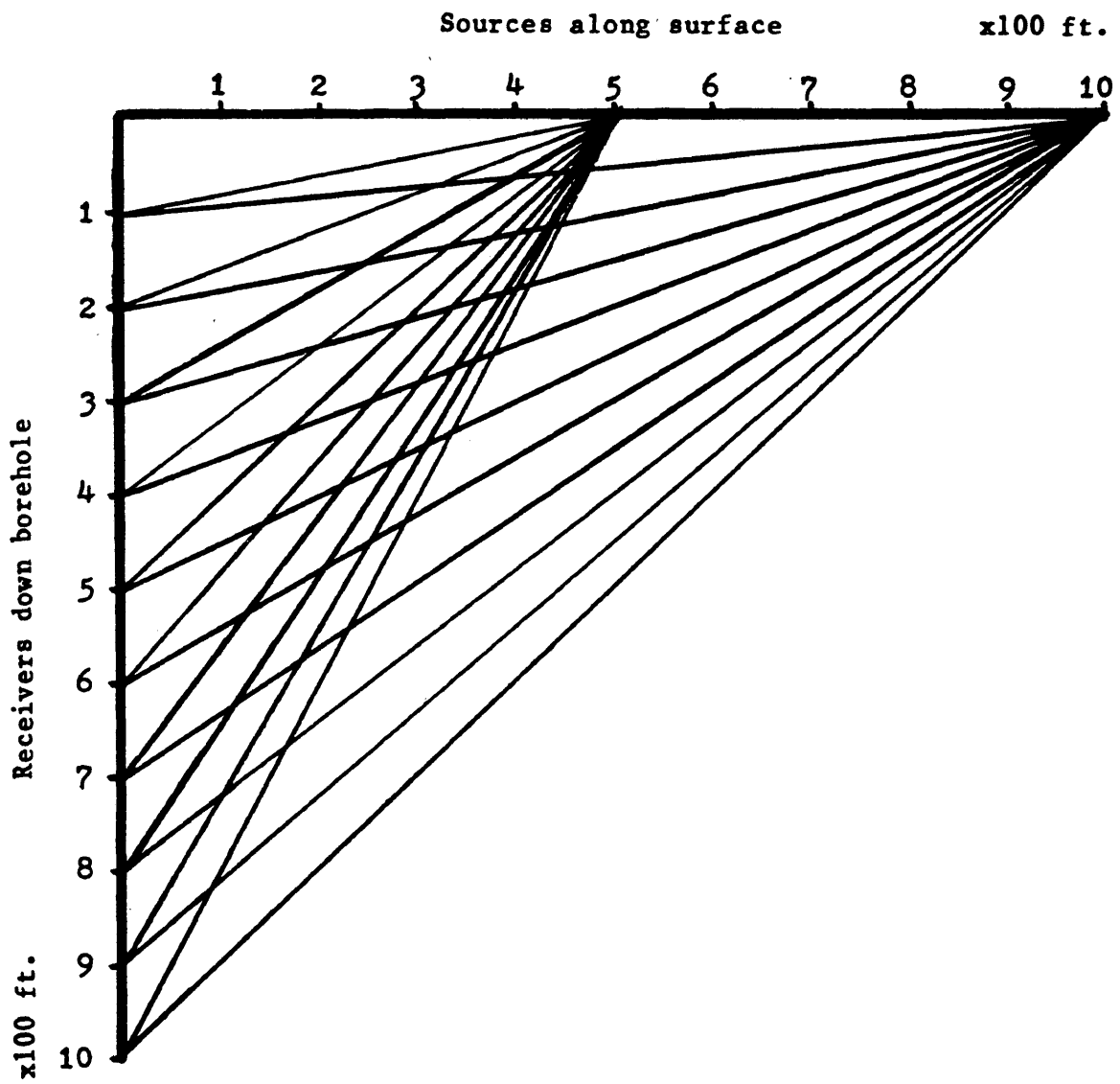


Figure 1. Multiple offset VSP geometry.

- 4) The average slowness at the center of a parameter circle is the difference of 2) and 3). These outputs form a matrix indexed by mn .
- 5) Using the generated slowness matrix from 4), the expected travel times are calculated.
- 6) The residual time is the observed minus the calculated travel time.
- 7) This residual error time is equally distributed along the entire length of the ray.
- 8) A slowness correction matrix is computed in the manner of steps 2), 3), and 4).
- 9) Steps 5) through 8) are iterated.

2.1 Constant background

As stated earlier, this investigation used two different velocity backgrounds for the medium. First the constant velocity case will be discussed before going into the more complicated linear $c(z)$ medium.

Using a constant velocity reference, the raypaths become straight lines from the i 'th source to the j 'th receiver. Denote the length of this ray as L_{ij} . Because the ray follows a straight course from $(x_i, 0)$ to $(0, z_j)$, the raypath is represented by the equation:

$$z = - \left[\frac{z_j}{x_i} \right] x + z_j. \quad (1)$$

Over the two-dimensional medium, a set of overlapping circles is drawn, whose centers form an uniform square grid (Fig. 2). These circles must completely cover the medium such that every point within the medium is contained in at least one of these parameter circles. The average slowness within any circle is assumed to be the slowness at the origin of the circle. These circles serve as a truncated least-distance averaging method which lies at the basis of Horn's algorithm (1978). Also, circles prove to be more convenient than a grid of squares.

Sources are positioned at the center of some, or all, of the circles along the surface and receivers are located at the origin of some, or all, of the circles at the borehole edge of the model. Clearly, the source or receiver spacing defines the largest possible grid size, or equivalently, the density of the parameter circles. Resolution is lost with too small of a density of parameter circles. Conversely, a greater grid density will increase the resolution only to the limit of the data.

Let the known value r denote the radius of any parameter circle. Then the circle, around the point (x_m, z_n) , is mathematically expressed as:

$$(x - x_m)^2 + (z - z_n)^2 = r^2. \quad (2)$$

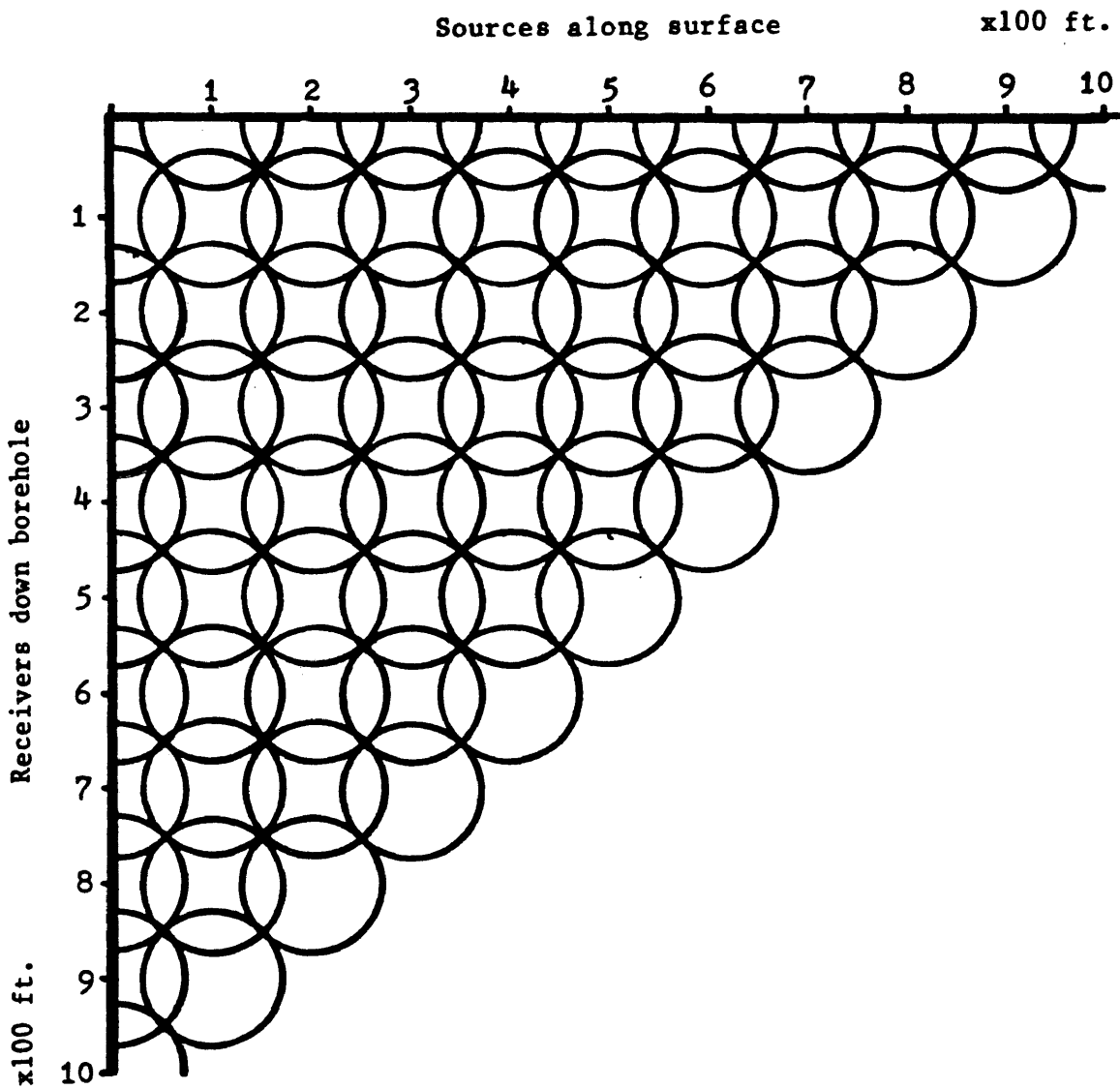


Figure 2. Modeling the medium using overlapping circles.

If the ray in question intersects the parameter circle, then equation (1) and equation (2) will have two points in common. A quadratic equation is formed by substituting (1) into (2). The quadratic is solved using the quadratic formula. Should both points prove to be imaginary, then the ray fails to intersect the circle.

The chord length l_{ij}^{mn} is calculated using the distance formula. These chord lengths are formed by all rays ij which intersect a given parameter circle around the point mn . The total length of ray segments d^{mn} , around the parameter mn , is the sum of all these chord lengths.

$$d^{mn} = \sum_{ij} l_{ij}^{mn}. \quad (3)$$

Every ray ij , which passes through a parameter circle mn , will have a residence time t_{ij}^{mn} within that circle as calculated from the observed travel time \tilde{T}_{ij} and a weighting factor w_{ij}^{mn} .

$$t_{ij}^{mn} = \tilde{T}_{ij} w_{ij}^{mn}. \quad (4)$$

This weighting factor depends on the background velocity and can be viewed as the ratio of the expected time spent inside the circle to the total expected time of the ray. For a constant c_0 velocity background:

$$w_{ij}^{mn} = \frac{\left[\frac{l_{ij}^{mn}}{c_0} \right]}{\left[\frac{L_{ij}}{c_0} \right]} = \frac{l_{ij}^{mn}}{L_{ij}}. \quad (5)$$

The total time spent about a parameter point mn is the sum of all the residence times. The average slowness at the point mn is given by:

$$s^{mn} = \frac{\sum_{ij} t_{ij}^{mn}}{\sum_{ij} l_{ij}^{mn}}. \quad (6)$$

2.2 Constant background - Iterative algorithm

An iterative technique is applied to improve upon this first guess model. First, a forward model is needed to compute an estimated travel time. Using straight raypaths, the ray is divided into equal segments Δl . The slowness along the entire segment length is assumed to equal the slowness at the midpoint of the segment, which is estimated by fitting a two-dimensional parabola to the surrounding four points of the slowness matrix. The estimated travel time T_{ij} is the sum of the slowness along the raypath.

$$T_{ij} = \sum_{(x_i, 0)}^{(0, z_j)} s(x, z) \Delta l. \quad (7)$$

The total residual travel time E_{ij} along ray ij is given by:

$$E_{ij} = \tilde{T}_{ij} - T_{ij}. \quad (8)$$

This residual error time is equally distributed along the length of the ray. The slowness correction matrix is the ratio of the total error times about a parameter to the total length of the ray segments.

$$\tilde{s}^{mn} = \frac{\sum_{ij} e_{ij}^{mn}}{\sum_{ij} l_{ij}^{mn}}, \quad \text{where } e_{ij}^{mn} = E_{ij} \left[\frac{l_{ij}^{mn}}{L_{ij}} \right]. \quad (9)$$

This slowness correction matrix is added to the previous slowness matrix to obtain an updated model. Iteration is performed by computing a new correction matrix, based on the current slowness model, until the mean square of the residual times reaches a steady state.

$$\sum_{ij} (E_{ij})^2_{\text{old}} - \sum_{ij} (E_{ij})^2_{\text{new}} < \epsilon. \quad (10)$$

2.3 Linear $c(z)$ medium

For a linear $c(z)$ medium, the raypaths are no longer straight, but instead are curved. The local radius of curvature is derived using Snell's Law.

$$\frac{\sin i}{c(z)} = p = \text{constant.} \quad (11)$$

Taking the derivative of the ray parameter, and noting that the derivative of any constant is zero, we find that

$$dp = \frac{\partial p}{\partial i} di + \frac{\partial p}{\partial z} dz, \quad (12)$$

or

$$0 = \frac{\cos i}{c(z)} di - \frac{\sin i}{c^2(z)} \frac{dc(z)}{dz} dz. \quad (13)$$

From the geometry shown in Fig. 3:

$$dS = R di, \quad \text{and} \quad \cos i = \frac{dz}{dS}. \quad (14)$$

By combining the two equations (14), we conclude that

$$di = \frac{dz}{R \cos i}. \quad (15)$$

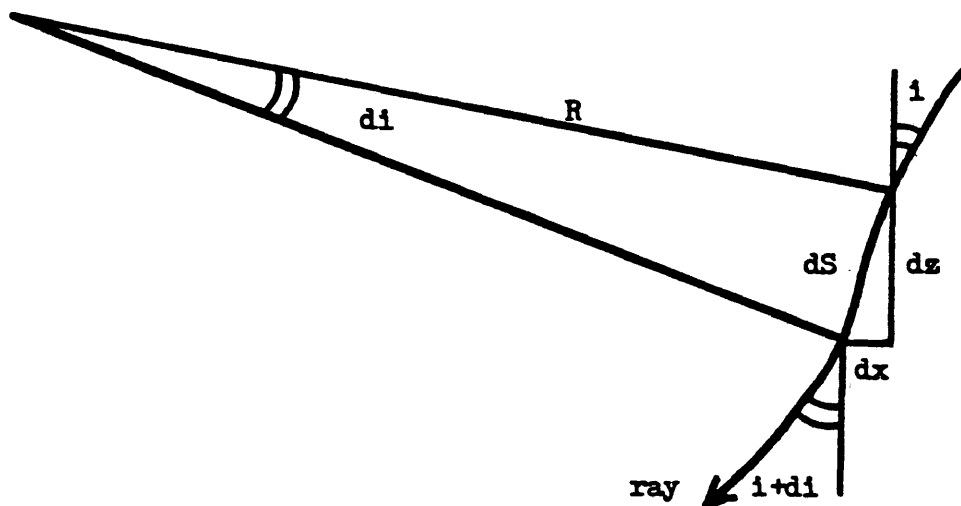


Figure 3. Derivation of the local radius of curvature.

Substituting (15) into (13):

$$0 = \frac{dz}{R c(z)} - \frac{\sin i}{c^2(z)} \frac{dc(z)}{dz} dz, \quad (16)$$

or

$$\frac{1}{R} = \frac{\sin i}{c(z)} \frac{dc(z)}{dz} = p \frac{dc(z)}{dz}. \quad (17)$$

So for a linear velocity $c(z) = c_0 + m_0 z$, the radius of curvature becomes a constant, which implies that the raypath is simply the arc along some circle; that is,

$$\frac{1}{R} = p m_0. \quad (18)$$

Using the Wiechert-Herglotz integral (Aki and Richards, 1980), the ray parameter can be determined (See appendix A). The result is

$$p = \frac{2 m_0 x_i}{\sqrt{[m_0^2 x_i^2 + c_0^2 + c^2(z_j)]^2 - 4 c_0^2 c^2(z_j)}}. \quad (19)$$

From the ray parameter, the radius of curvature becomes known using equation (18). With two points on this circular path known, namely the shot and receiver locations, the ray follows a path on the circle:

$$(x - x_0)^2 + (z - z_0)^2 = R^2. \quad (20)$$

Using the geometry in Fig. 4, the center of the circle can be determined:

$$|x_0| = R \cos(i + \alpha); \quad (21)$$

$$|z_0| = R \sin i. \quad (22)$$

Both angles are calculated from the geometry. The results are

$$\sin \frac{\alpha}{2} = \frac{x_i^2 + z_j^2}{2R}; \quad (23)$$

$$\tan \left(i + \frac{\alpha}{2}\right) = \frac{x_i}{z_j}. \quad (24)$$

Thus for a linear $c(z)$ medium, the rays are arcs of a circle. The overall length of the ray is:

$$L_{ij} = R\alpha. \quad (25)$$

Over the two-dimensional medium, a grid of overlapping circles is

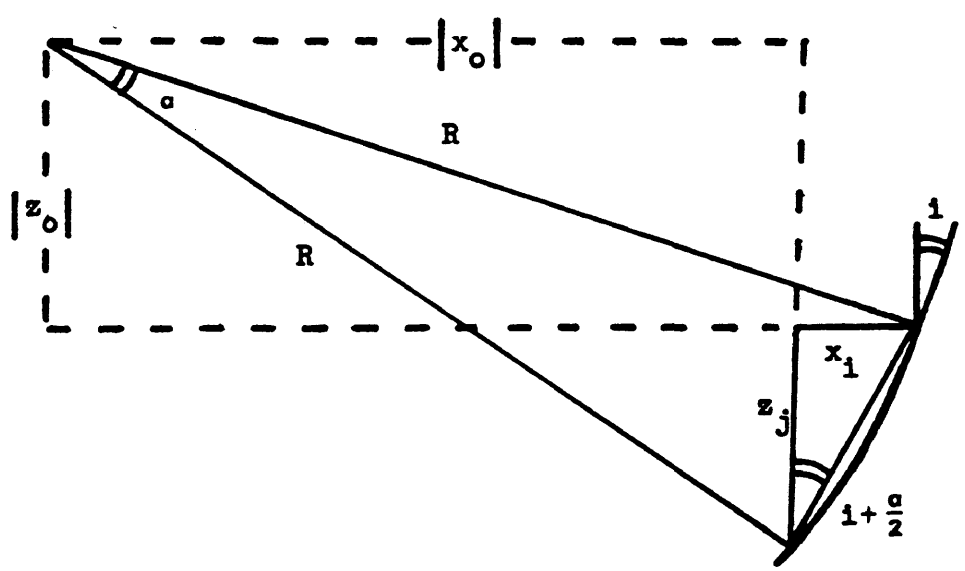


Figure 4. Origin determination.

superimposed. Any ray will intersect a circle provided that:

$$|d - R| < r, \quad \text{where } d^2 = (x_m - x_o)^2 + (z_n - z_o)^2. \quad (26)$$

The arc length l_{ij}^{mn} formed by the ray ij , within the circle mn , is computed using the law of cosines (Fig. 5).

$$\cos \frac{\beta}{2} = \frac{R^2 + d^2 - r^2}{2 R d}; \quad (27)$$

$$l_{ij}^{mn} = R \beta. \quad (28)$$

The total ray distance d^{mn} , about a given parameter, is the sum of all the arc lengths.

$$d^{mn} = \sum_{ij} l_{ij}^{mn}. \quad (29)$$

Every ray ij , which passes through a parameter circle mn , will have a residence time t_{ij}^{mn} within that circle as calculated from the observed travel time \tilde{T}_{ij} and a weighting factor w_{ij}^{mn} .

$$t_{ij}^{mn} = \tilde{T}_{ij} w_{ij}^{mn}. \quad (30)$$

The weighting factor is the ratio of the expected time spent inside

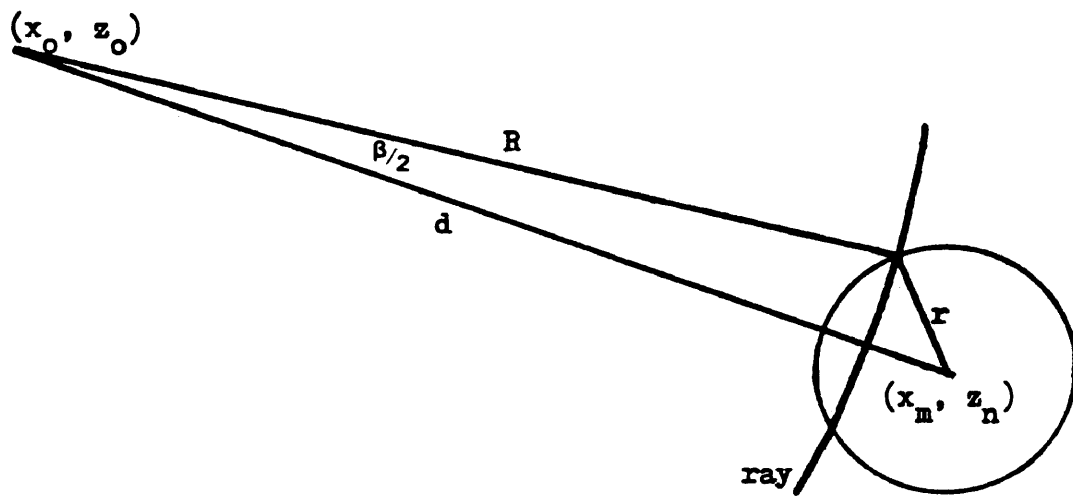


Figure 5. Arc length determination.

the parameter circle to the overall travel time for the entire ray.

$$w_{ij}^{mn} = \frac{\left[\frac{l_{ij}^{mn}}{c(z_n)} \right]}{T_{ij}}. \quad (31)$$

Here $c(z_n)$ is the velocity at the origin of the parameter circle and T_{ij} is the theoretical travel time of the ray in the assumed linear background (see appendix B).

$$T_{ij} = \frac{1}{m_o} \ln \left[\frac{c(z_j) (1 + \sqrt{1 - p^2 c_o^2})}{c_o (1 + \sqrt{1 - p^2 c^2(z_j)})} \right]. \quad (32)$$

The average slowness at the point mn is given by:

$$s_{ij}^{mn} = \frac{\sum_{ij} t_{ij}^{mn}}{\sum_{ij} l_{ij}^{mn}}. \quad (33)$$

2.4 Linear $c(z)$ medium - Iterative algorithm

The slowness matrix can be improved by an iteration technique. The curved ray is divided into equal segments Δl . The slowness along the entire segment length is assumed to equal the slowness at the midpoint of the segment, which is estimated by fitting a two-dimensional para-

boloid to the surrounding four points of the slowness matrix. An estimated travel time is computed by summing the slownesses at the mid-points of the segments.

$$T_{ij} = \int_{(x_i, 0)}^{(0, z_j)} s(x, z) \Delta l. \quad (34)$$

The total residual time E_{ij} along the ray ij is given by:

$$E_{ij} = \bar{T}_{ij} - T_{ij}. \quad (35)$$

This residual error time is distributed equally along the raypath. The slowness correction matrix is the ratio of the sum of the local residual times to the sum of the arc lengths around each parameter:

$$\bar{s}^{mn} = \frac{\sum_{ij} e_{ij}^{mn}}{\sum_{ij} l_{ij}^{mn}}, \quad \text{where } e_{ij}^{mn} = E_{ij} \left[\frac{l_{ij}^{mn}}{L_{ij}} \right]. \quad (36)$$

This slowness correction matrix is added to the slowness matrix to obtain a revised slowness model. Iteration is performed, by computing a new correction matrix based on the updated slowness model, until the mean square of the residual times reaches a steady state:

$$\sum_{ij} (E_{ij})^2_{old} - \sum_{ij} (E_{ij})^2_{new} < \epsilon. \quad (37)$$

3. TESTING THE ALGORITHMS

Both the constant and linear velocity background ART algorithms were tested using synthetic data sets. For a given slowness model, the travel times were generated using the forward modeling scheme in either the constant or linear velocity case. The slowness was integrated along either straight or circular paths as dictated by the general velocity trend of the input velocity model. Where the velocity increased linearly with x instead of z , the reciprocity principle of interchanging the shot and receiver locations was applied. Such data was calculated by interchanging the x and z variables and using the technique of section 2.4 to compute the synthetic data.

Clearly, the synthetic data does not correspond exactly to the perturbations which would be measured in the real case, but instead, the travel times represent a first order approximation for relatively small amplitude velocity anomalies. Therefore, my results test the accuracy of the reconstruction method and not the validity of linearizing the

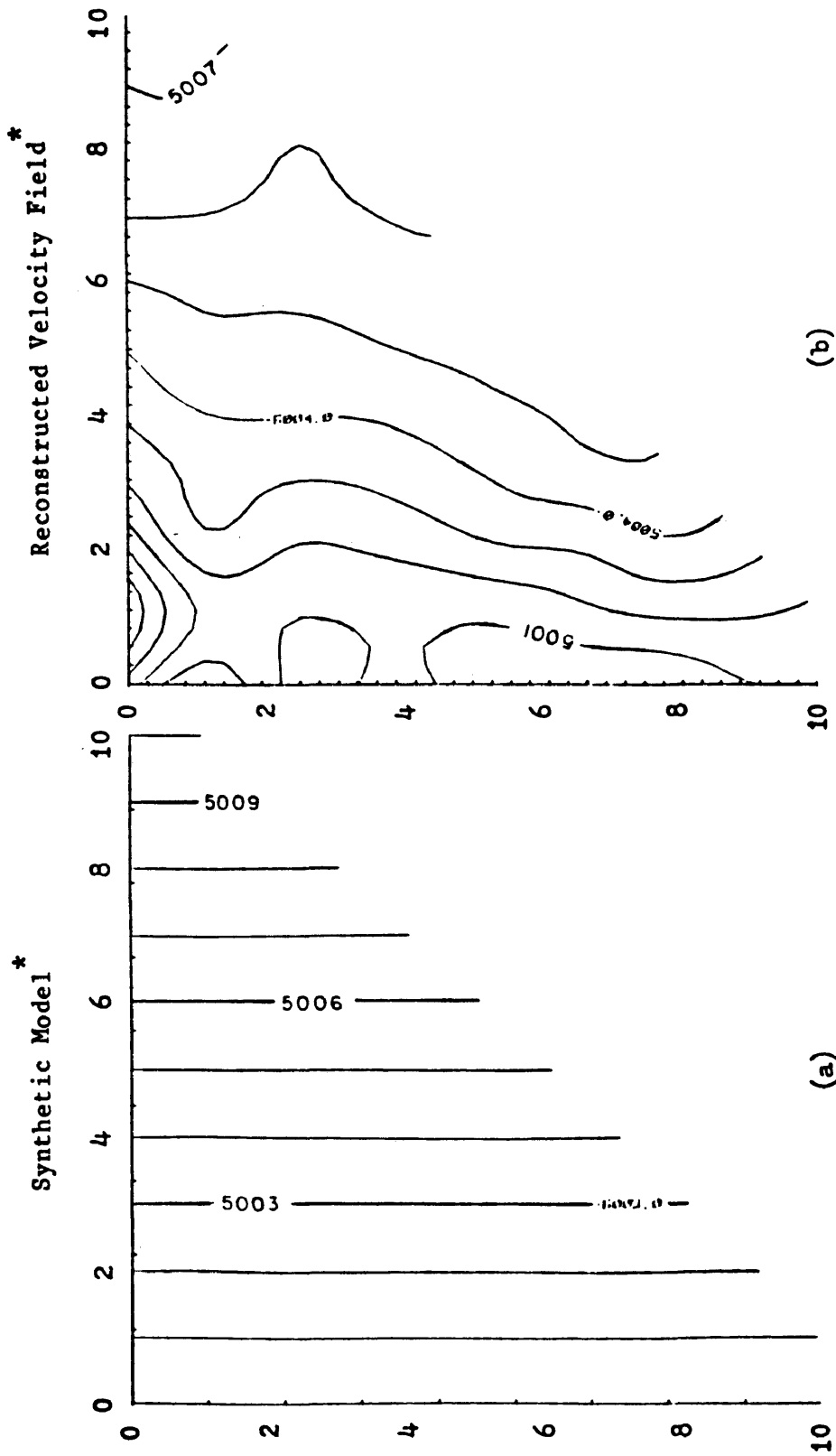
problem.

4. RESULTS

ART produces the best tomographical results when the velocity function is smoothly varying. Using synthetic data sets, accurately inverted velocity values occurred within the model, but the velocity values varied from the synthetic input near the edges of the model (Fig. 6, 7, 8, 9, 10, 11).

The greatest deviation occurred near the origin (Fig 6b, 7b, 8b, 11b, 11c). The first velocity down the well is erroneously high while the first velocity along the surface is erroneously low. The apparent cause of this phenomenon rests within the raypaths themselves and the way in which the ART algorithm handles these raypaths.

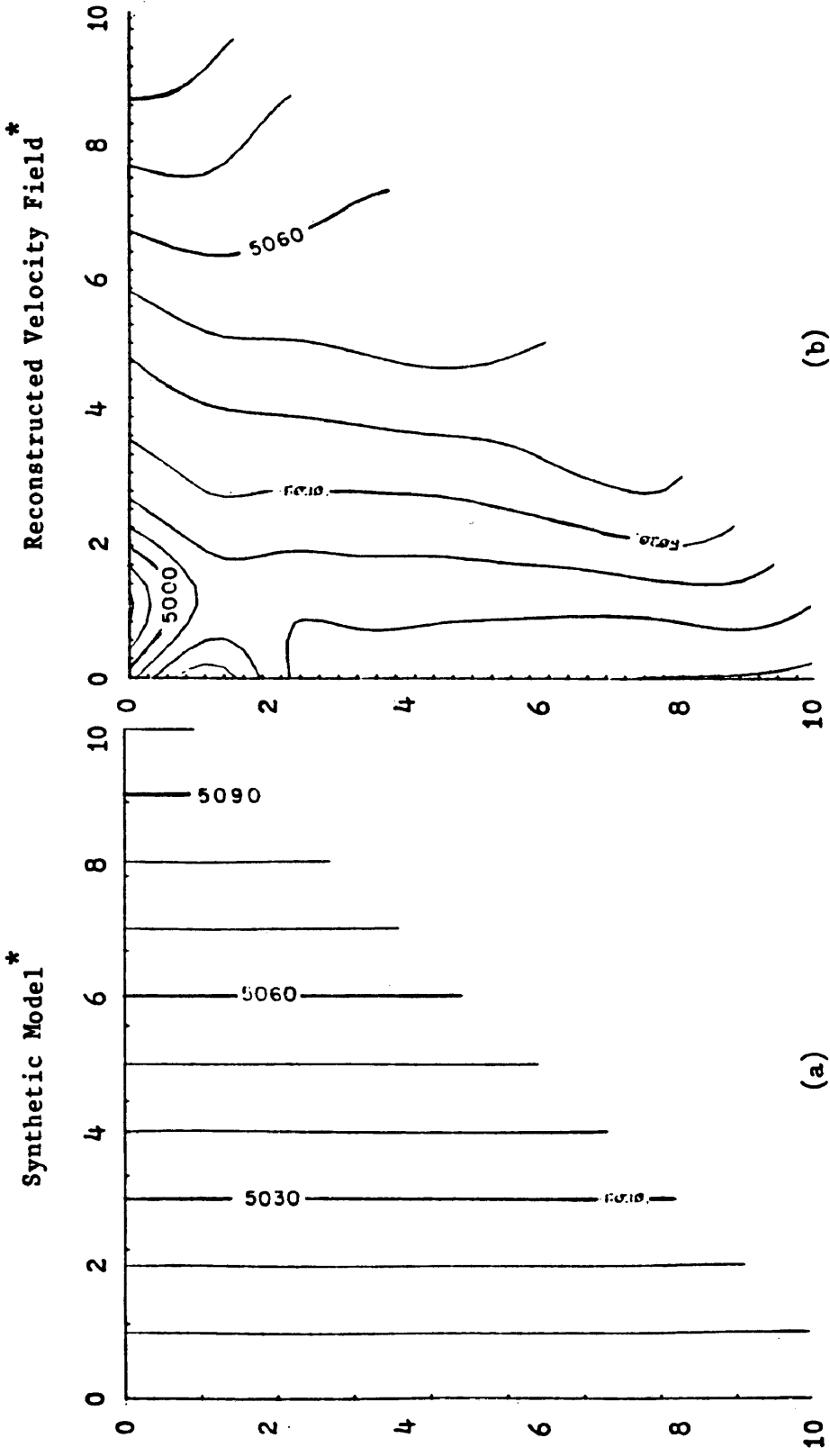
The first parameter down the well reflects the slownesses along the rays from all ten shot locations. However, the shortest raypath length most accurately reflects the true velocity at this point, while the longer rays reflect less of the velocity in question and more of the velocities further away. Because the ART algorithm equally weights all rays, regardless of their relative lengths, the slowness down the hole reflects the velocity change across the surface of the model. Appropriate constraints down the well will eliminate this problem, since the slowness in a borehole are usually known from the sonic log (Fig 11d).



* Travel times calculated using curved rays based on a background velocity $-v(x) = 5000 + .01(x_i)$.

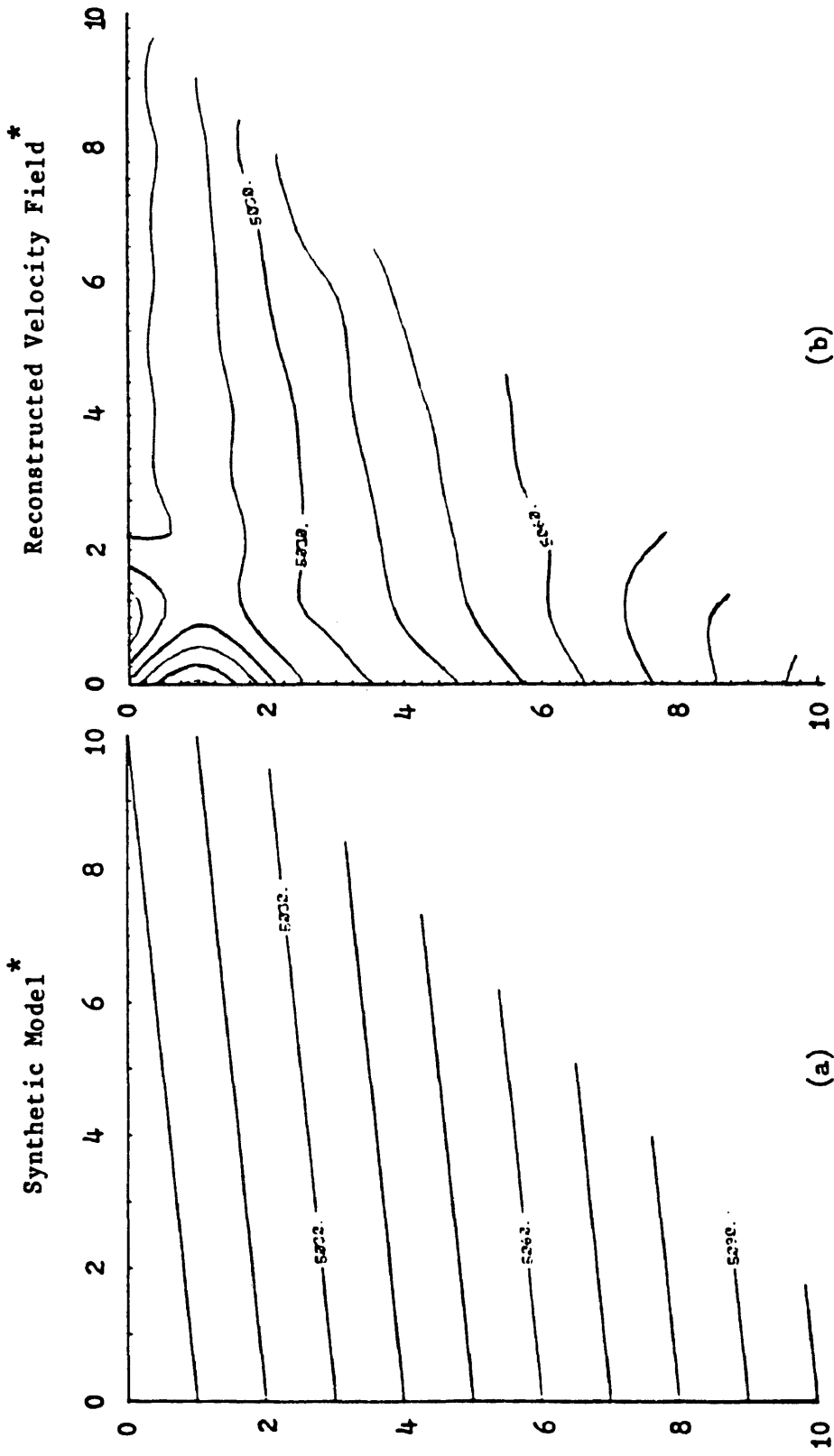
* After 7 iterations using straight rays.

Figure 6. Velocity increasing with x.



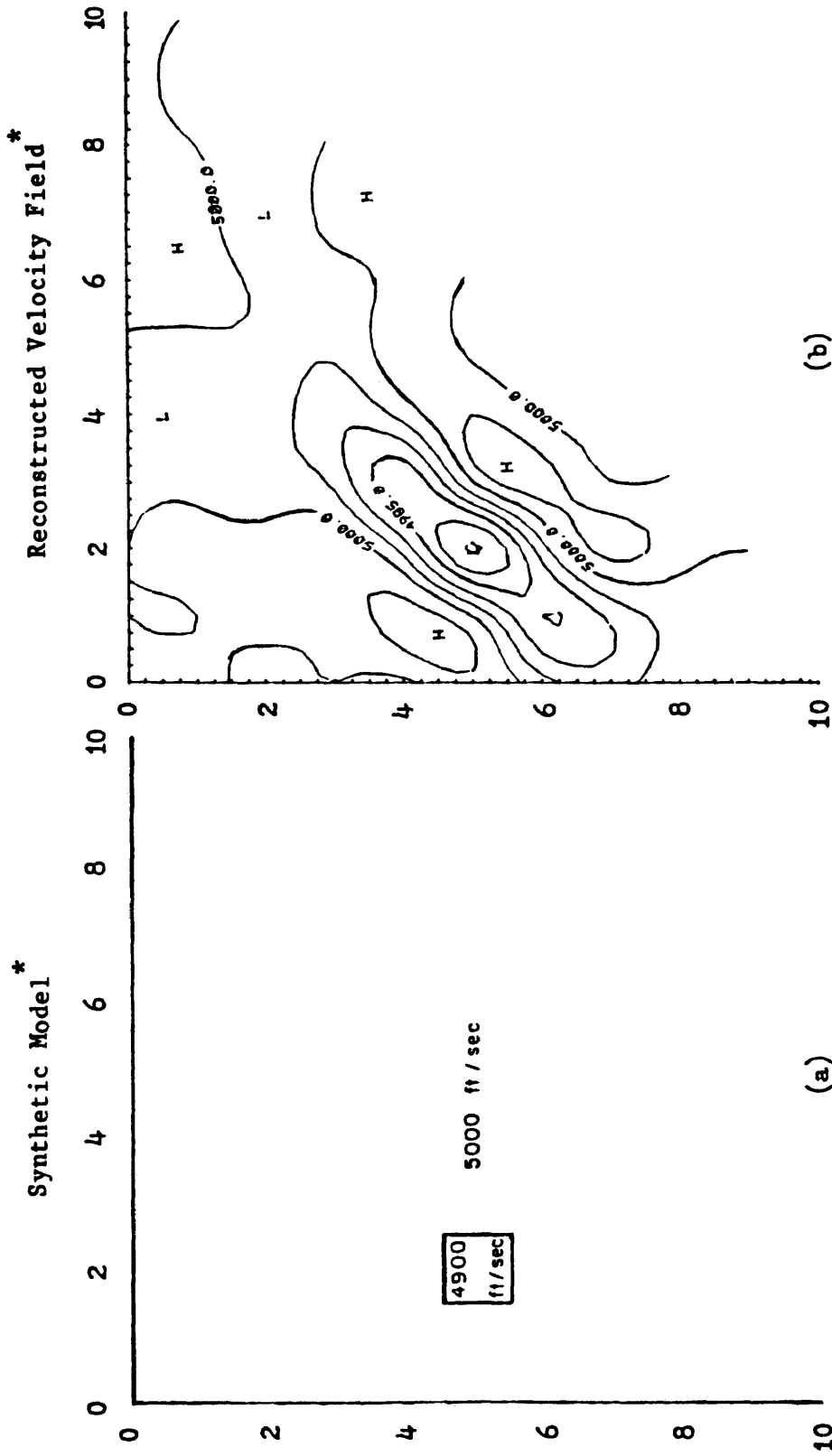
* Travel times calculated using curved rays based on a background velocity $- v(x) = 5000 + .1(x_i)$.
* After 9 iterations using straight rays.

Figure 7. Velocity increasing with x.



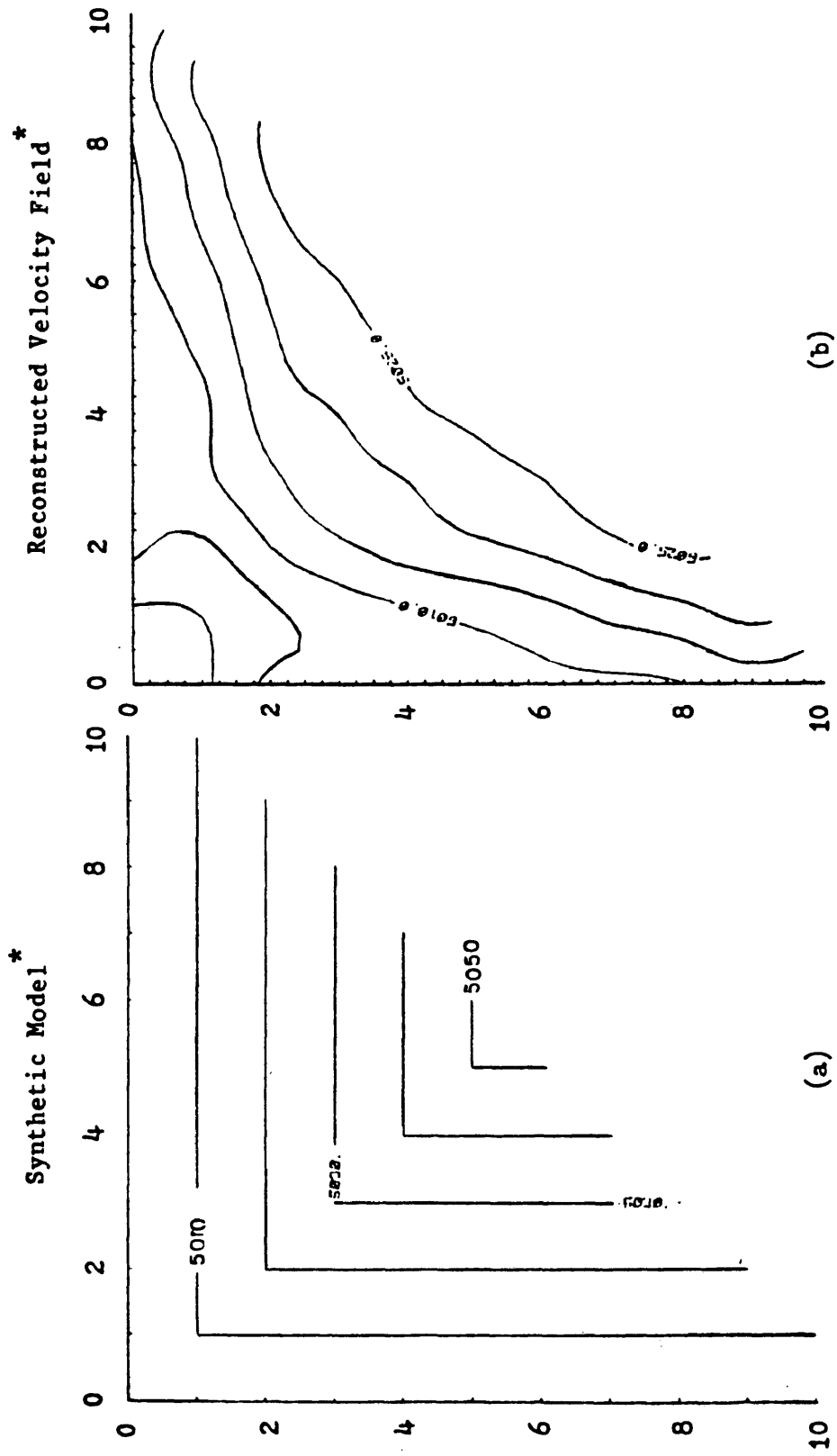
* Travel times calculated using straight rays. * After 11 iterations using straight rays.

Figure 8. Dipping model.



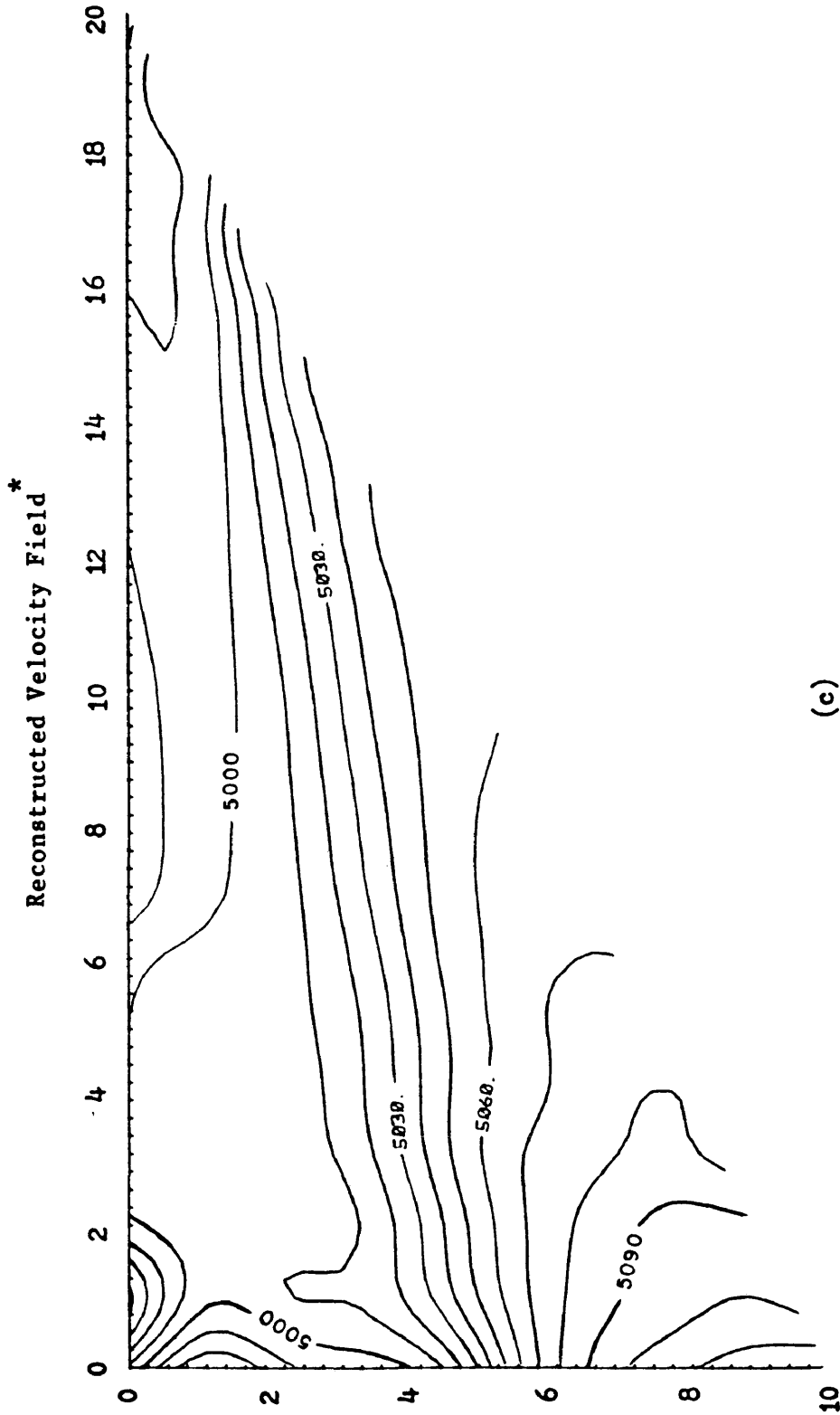
* Travel times calculated using straight rays. * After 12 iterations using straight rays.

Figure 9. Model of a low velocity anomaly.



* Travel times calculated using straight rays. ** After 7 iterations using straight rays.

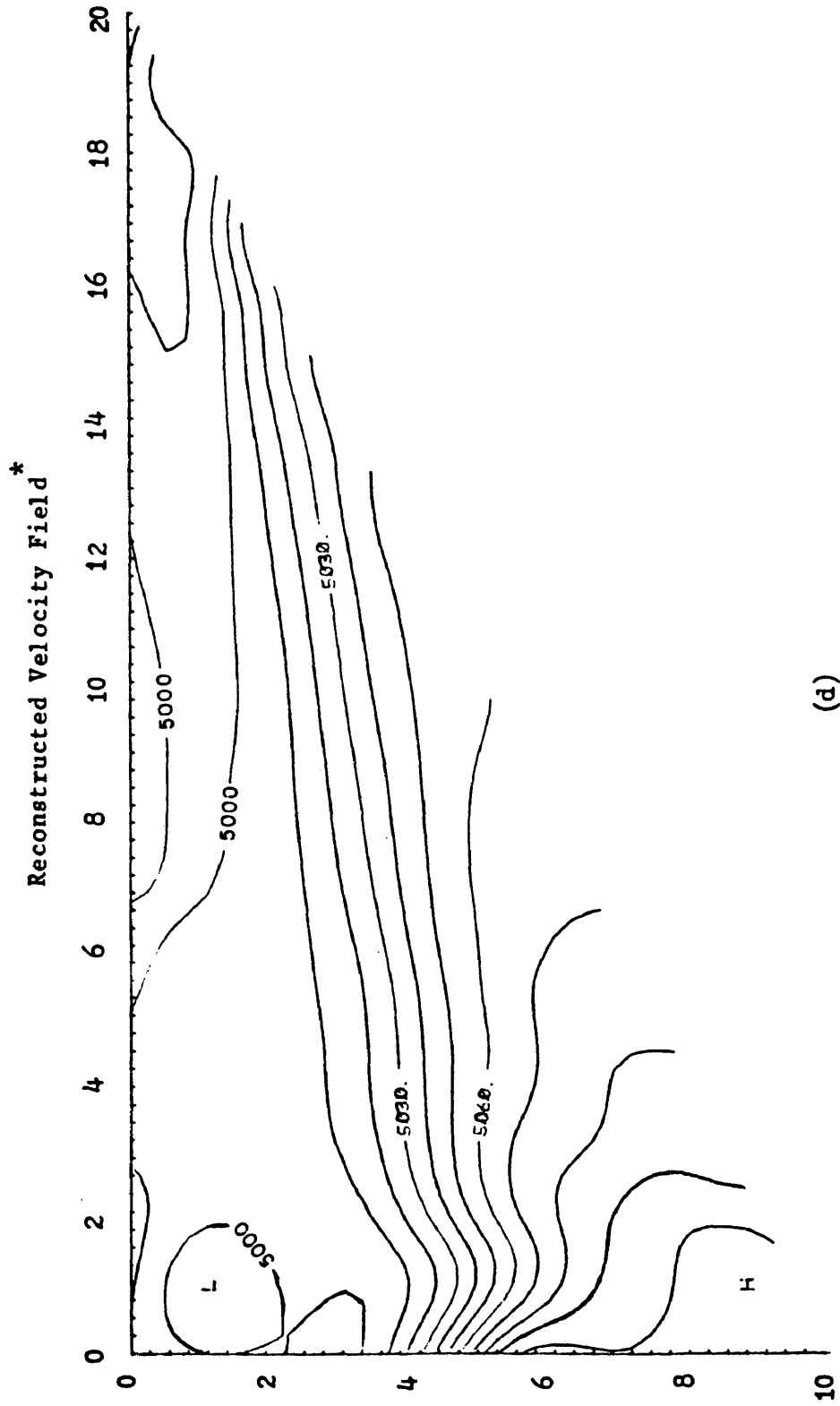
Figure 10. Velocity increasing with x and z.



(c)

* After 10 iterations using straight rays.

Figure 11. Two-layered model.



(d)

* After 12 iterations using straight rays. Velocity parameters were constrained down the well (z-axis).

Figure 11. Two-layered model.

Along the edges, the velocity values fluctuated from the synthetic model. These deviations resulted from two factors. First, the erroneous origin velocities greatly influence the surface and borehole edges. Second, the diagonal edge parameters are poorly constrained.

Unconstrained parameters are defined by only one ray. All parameters within my model had at least three rays defining each parameter. However, along the diagonal edge, these three rays are nearly identical and as such, they are defined by nearly the same parameters. Therefore, these rays fail to be completely independent of each other. Thus, the parameters along the diagonal edge are poorly constrained. In unconstrained cases, ART tends to average the velocities equally along the raypath resulting in resolution problems.

By using a delta function as the input velocity structure (Fig 9), we can better illustrate this resolution problem. Resolution along the raypaths is poor, but between the rays, the resolution increases. This accounts for the skewness of the inverted delta function into an elliptical form elongated along the dominant raypath direction. Still, the location of the slowness delta point is accurately defined, though not its true velocity value.

This poor resolution along the raypaths contributes to uniqueness problems (Fig 10). In general, an infinite number of models can produce the same travel time data. ART converges to the model which is the closest to the assumed background velocity and not necessarily a

representative of the true medium.

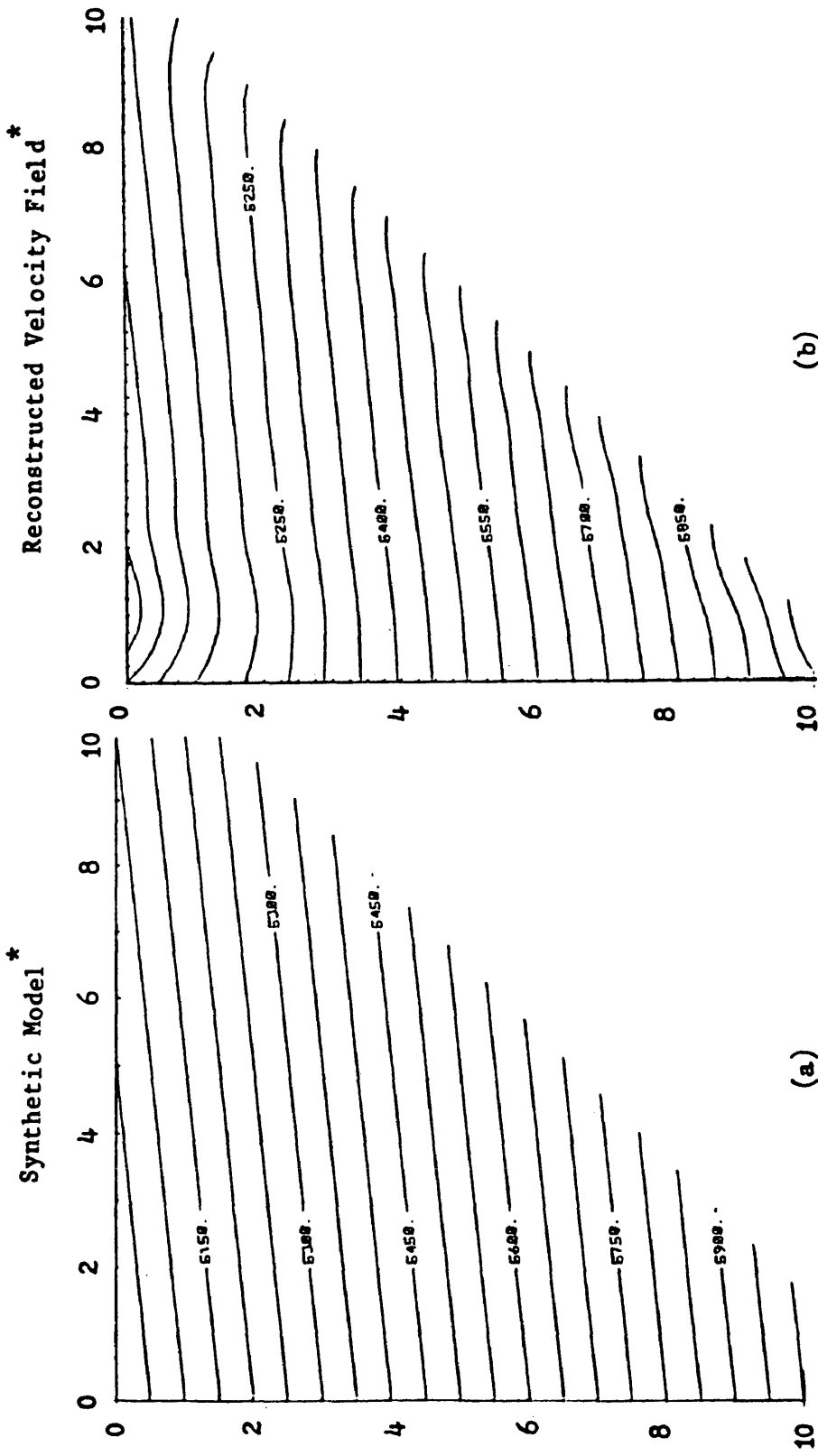
Uniqueness requires that all the slowness parameters are over constrained; that is more than one ray intersects each parameter circle. Furthermore, these rays must be independent. A close examination of my model reveals that at least three rays intersect each parameter; however, these rays travel approximately along the same direction and thus, these rays reflect many of the same slowness parameters. Clearly, these rays are not independent of each other.

The back-projected algorithm along curved rays was tested using synthetic data generated on linear trending, velocity models. In each case, the curved ART algorithm accurately reconstructed the synthetic models (Fig. 12, 13, 14).

5. PIERRE SHALE EXPERIMENT

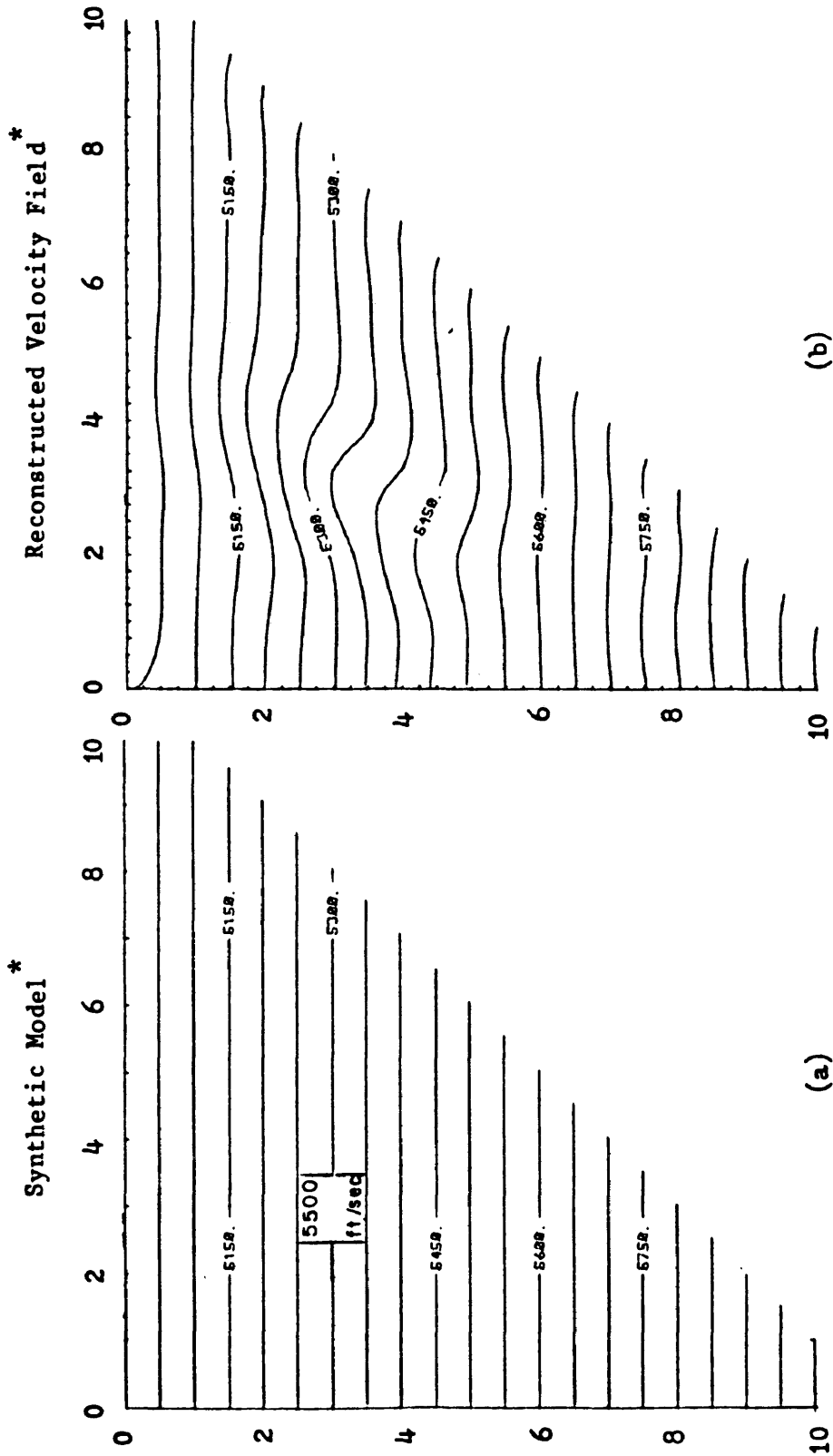
The multiple offset VSP, which was taken by the CSM Exploration Research Laboratory of the Geophysics Department, provided a real data test of the algorithms. Field work was done, in November 1980, at the CSM test site in northeastern Colorado, about five miles south of Brush Colorado (Southwest 1/4, Section 28, Township 3N, Range 55W). Sixta (1982) describes the data acquisition and data processing of this VSP experiment.

This site was chosen for its simple geologic setting. A 40 foot



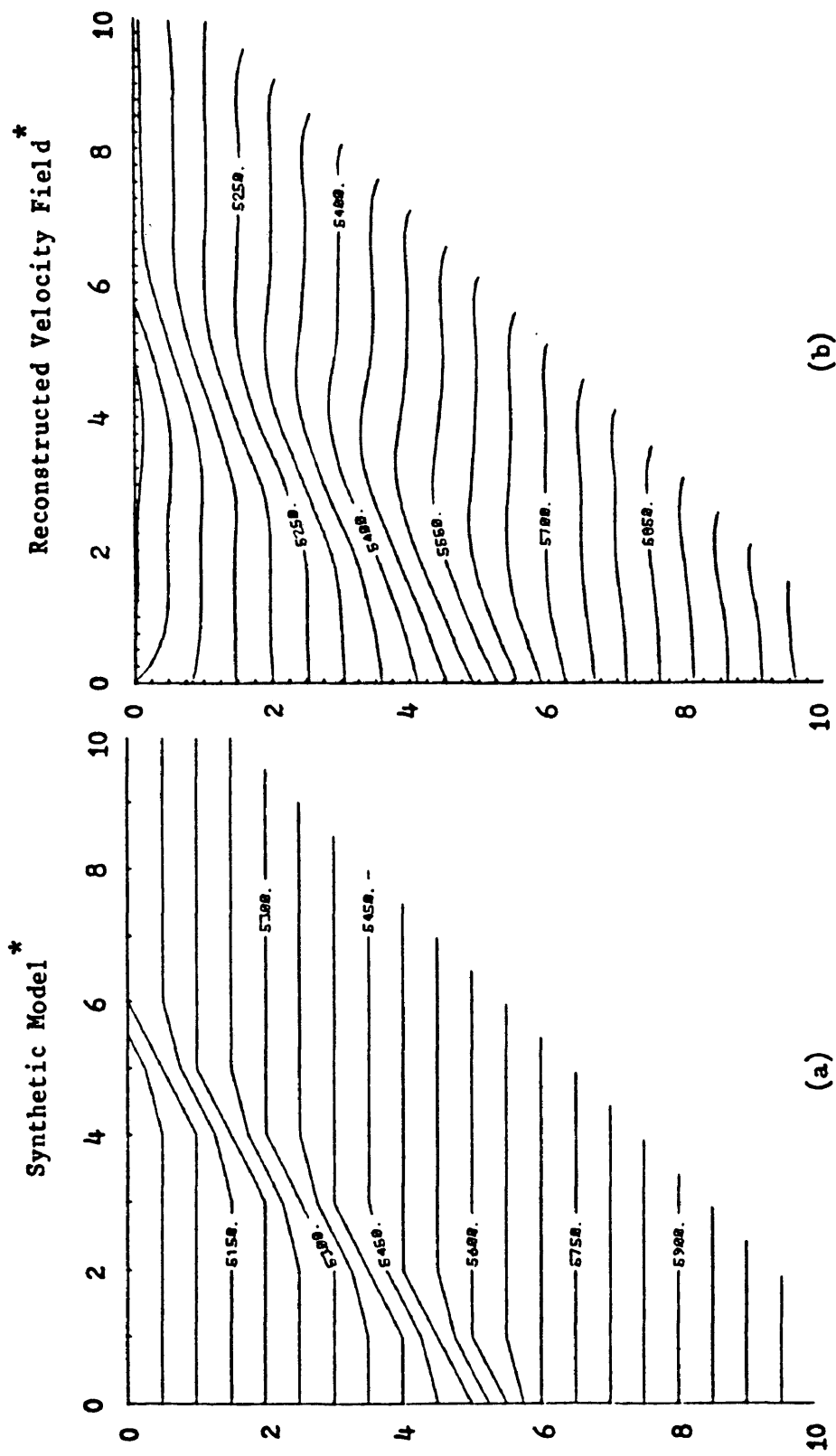
* Travel times calculated using curved rays based on a background velocity $- v(z) = 5000 + z_j$.
* After 10 iterations using curved rays.

Figure 12. Dipping model.



* Travel times calculated using curved rays based on a background velocity $- v(z) = 5000 + z_j$.
* After 10 iterations using curved rays.

Figure 13. High velocity anomaly within a linear trend.



*Travel times calculated using curved rays based * After 10 iterations using curved rays.
on a background velocity - $v(z) = 5000 + z_j$ ft./sec.

Figure 14. Faulted model.

thick layer of eolian sand rests on the surface. Beneath this sand layer, a 40 foot thick layer of clay grades into the Pierre Shale below. These thicknesses vary laterally away from the borehole. Because the upper 1000 feet of the Pierre Shale is well noted for its homogeneous nature, refracted events should not appear on the records and therefore, the direct ray will be the first arrival.

The field layout is sketched in Fig. 15. Geophones were spaced down the cemented well every 100 feet within the interval between 400 and 1000 feet. Dynamite, placed in shot holes at a depth of 100 feet deep, were shot at 200 foot intervals across the surface to a distance of 1400 feet from the borehole. The travel times were picked at the first trough on the records (Fig. 16) and are tabulated in the table below.

Table 1: Travel times in msec.

Depth in ft.	Horizontal offset in ft.						
	200	400	600	800	1000	1200	1400
400	50.25	62.00	82.75	106.25	133.50	160.00	187.00
500	64.50	74.00	91.75	113.00	138.50	163.50	189.75
600	78.75	86.50	101.75	121.00	144.50	168.00	193.50
700	92.25	99.00	112.25	129.75	151.50	174.00	198.25
800	106.00	111.75	123.50	139.25	159.50	180.75	203.50
900	119.25	124.50	135.00	149.00	168.00	188.00	210.00
1000	133.50	138.00	147.00	160.00	177.75	196.50	217.00

Picking the arrival times at the first trough results in a lag time of about 3 msec. from the true onset of the waveform. This error in the travel time corresponds to a similar error in the velocity value. The

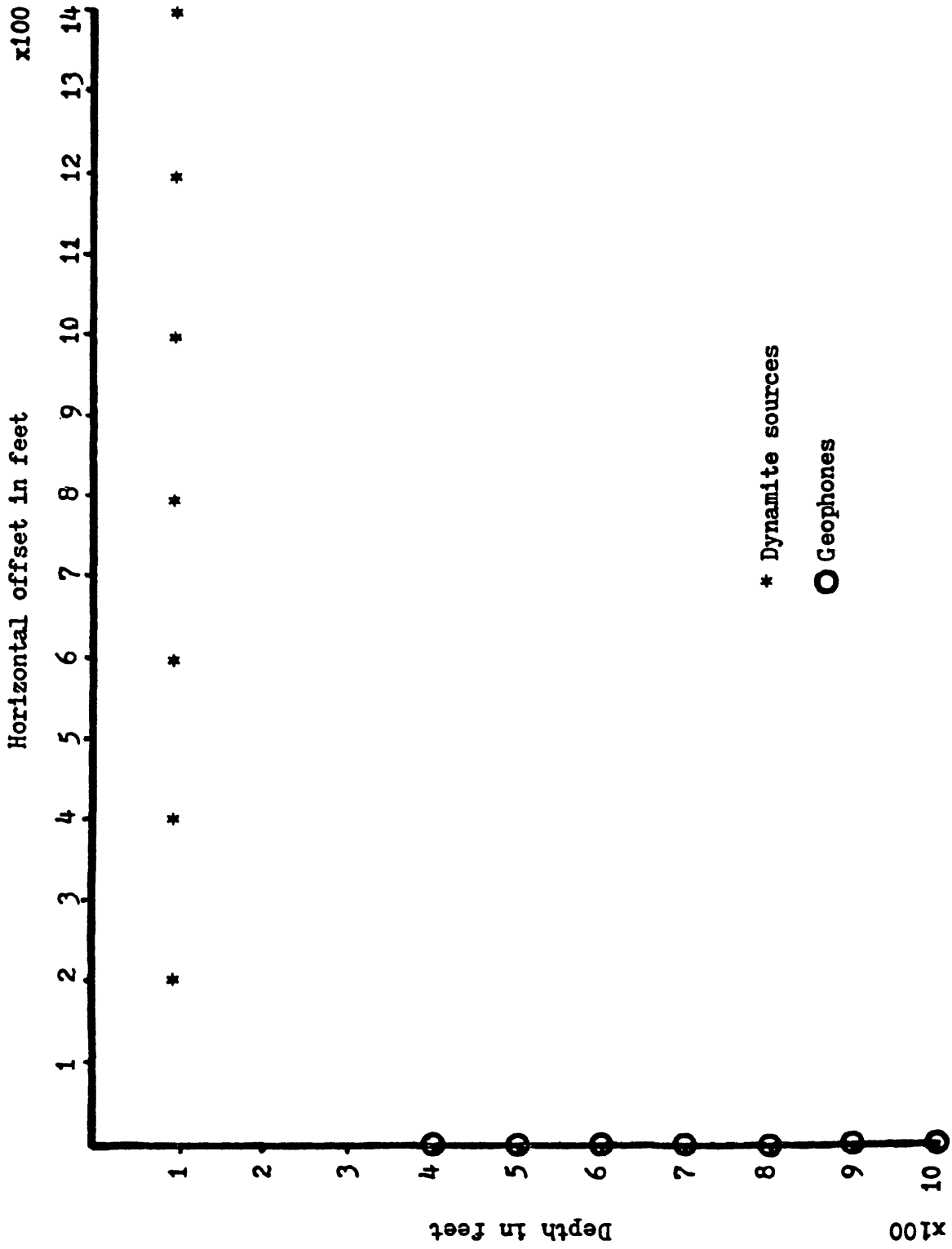


Figure 15. Field experiment layout.

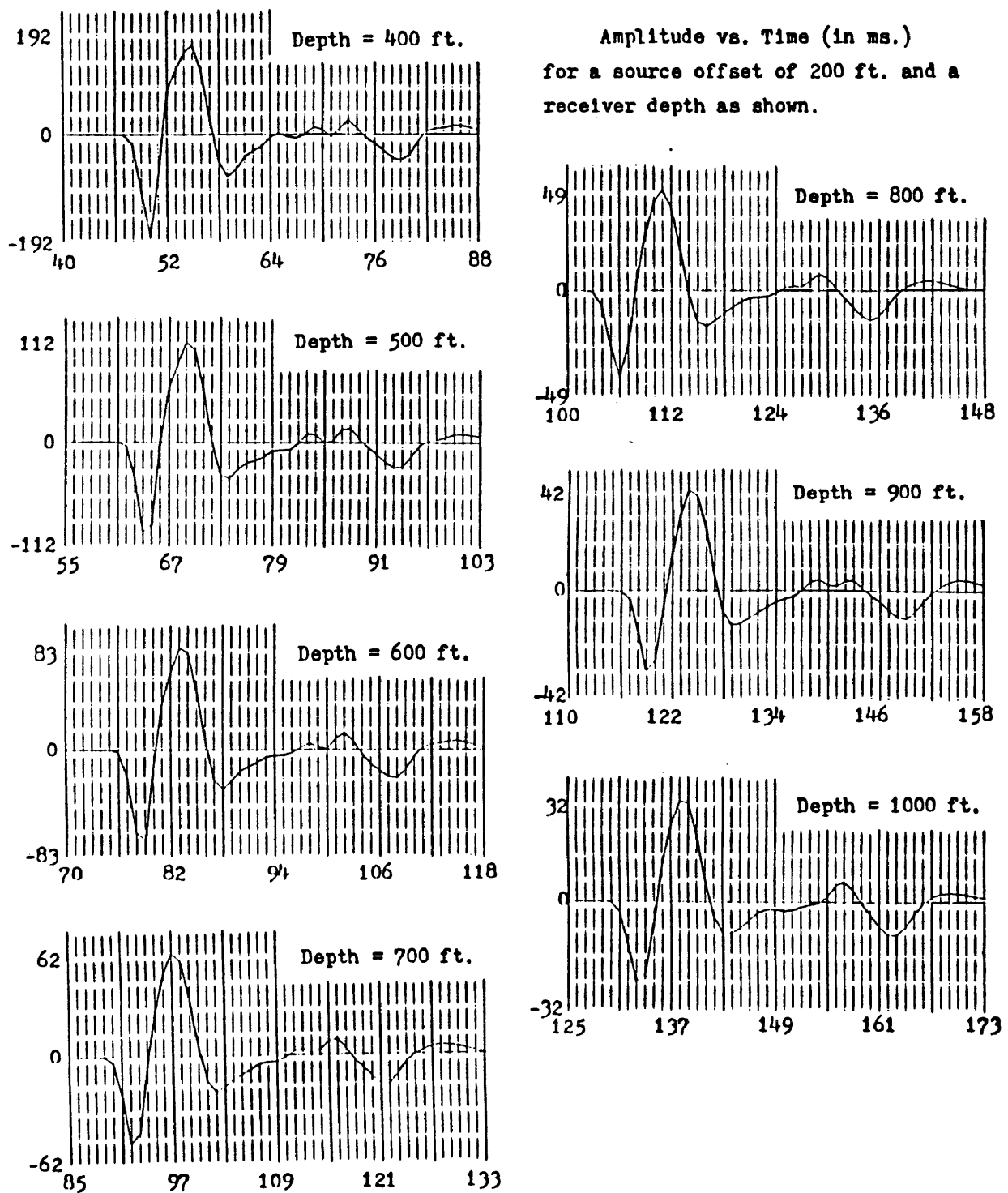


Figure 16. Typical waveforms.

travel time is the integral of the slowness along the raypath.

$$t = \int \frac{1}{c} dL, \quad \text{or} \quad t \approx \frac{L}{c}. \quad (38)$$

Differentiating (38), we obtain

$$dt \approx -L \frac{dc}{c^2}. \quad (39)$$

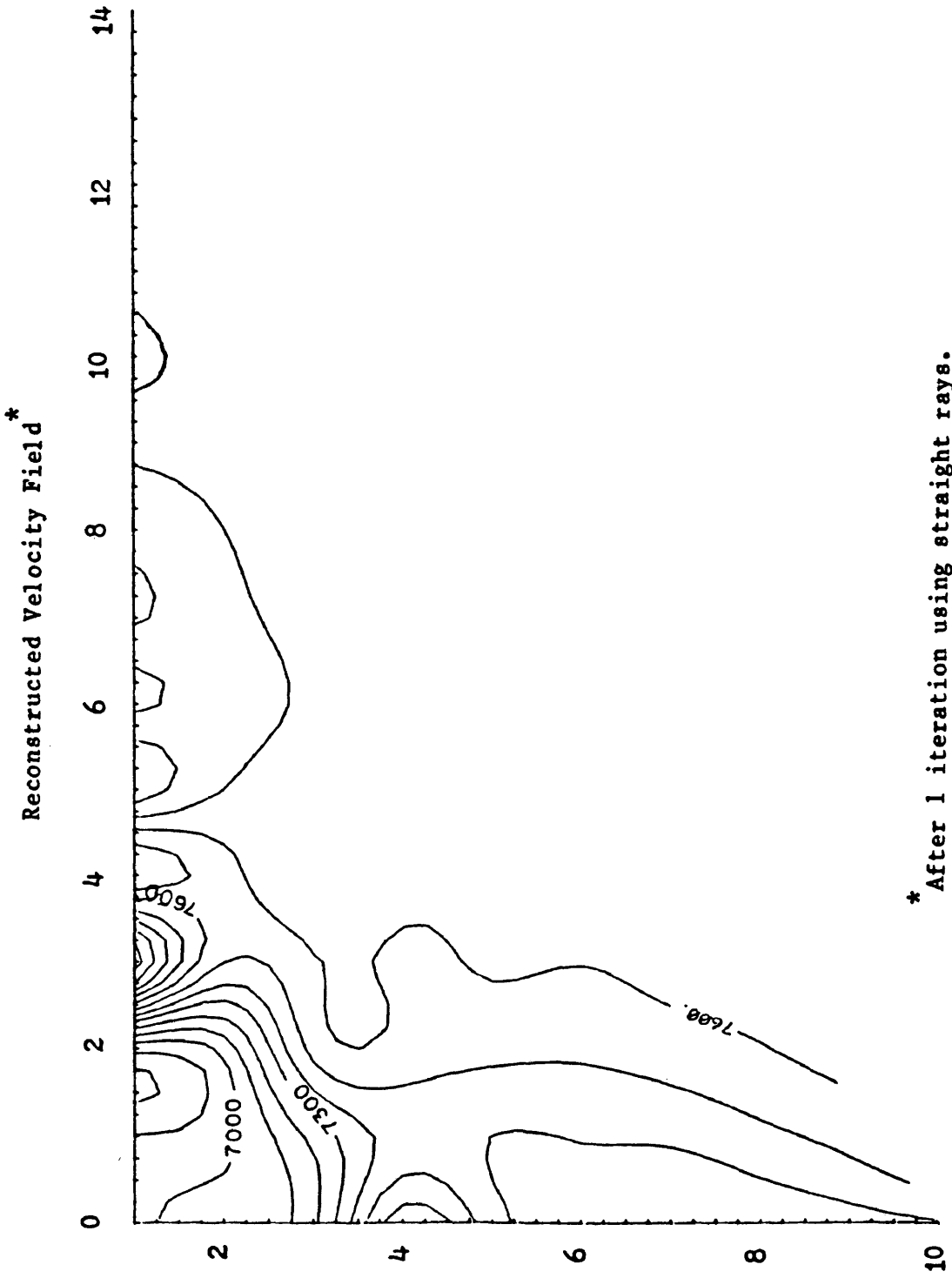
Dividing (38) into (39), yields the result

$$\frac{dt}{t} \approx -L \frac{dc}{c}. \quad (40)$$

By setting $dt=3$ msec. and noting that t ranges from 50.25 to 217.0 msec., the true velocity values range from 6.0 to 1.3% faster than those which are shown in the reconstructed fields.

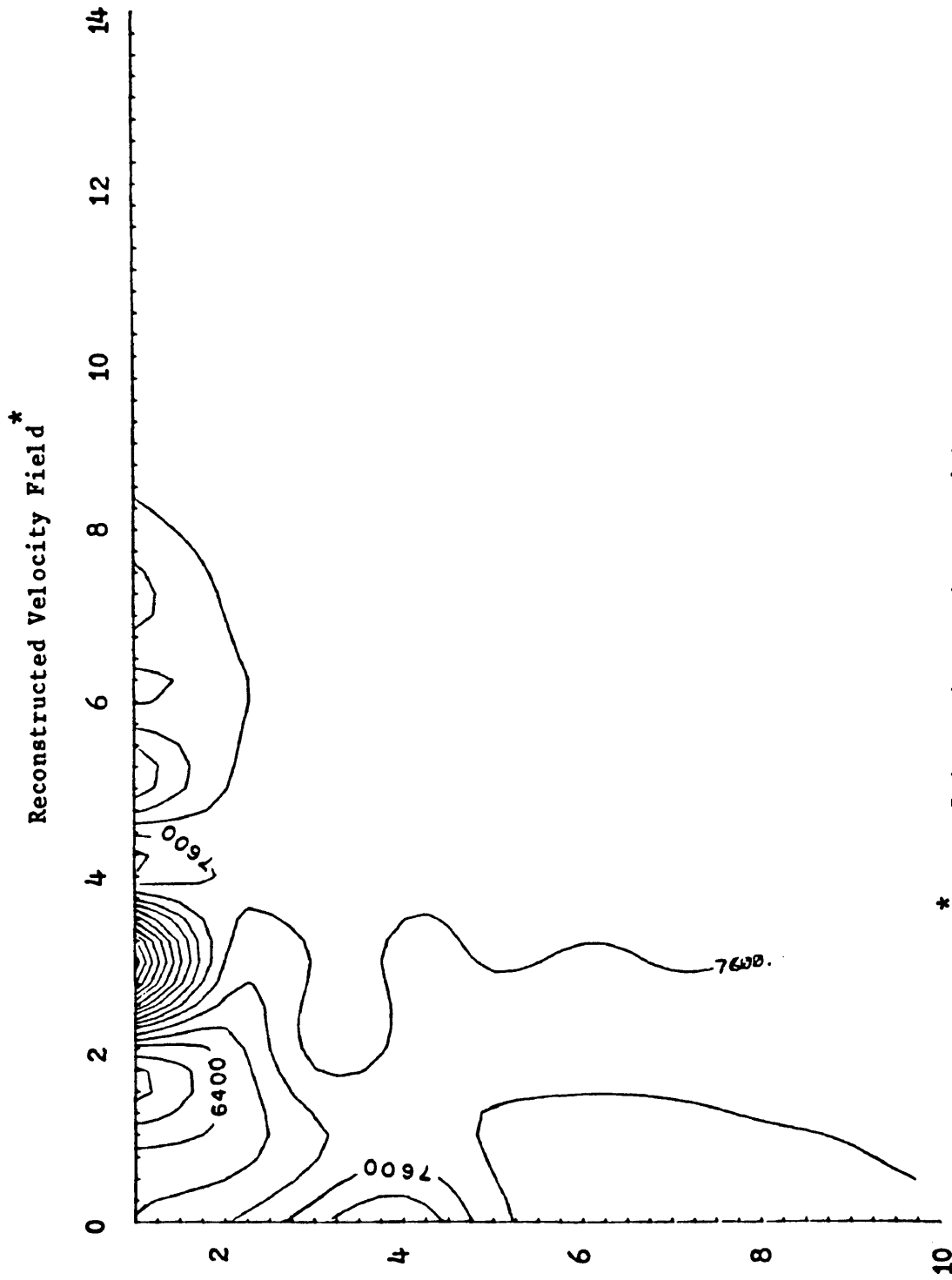
The Pierre Shale is well noted for its homogeneous nature; however, our reconstructions fail to support this characteristic. Imaging was performed both without parameter constraints and with constraining the slowness parameters down the borehole to within five percent of the measured slowness values off the sonic log.

Both the constrained (Fig. 18a, 18b) and the unconstrained (Fig. 17a, 17b) reconstructions reveal a questionable high velocity cell near the surface. Although the sonic log for the well records a velocity of



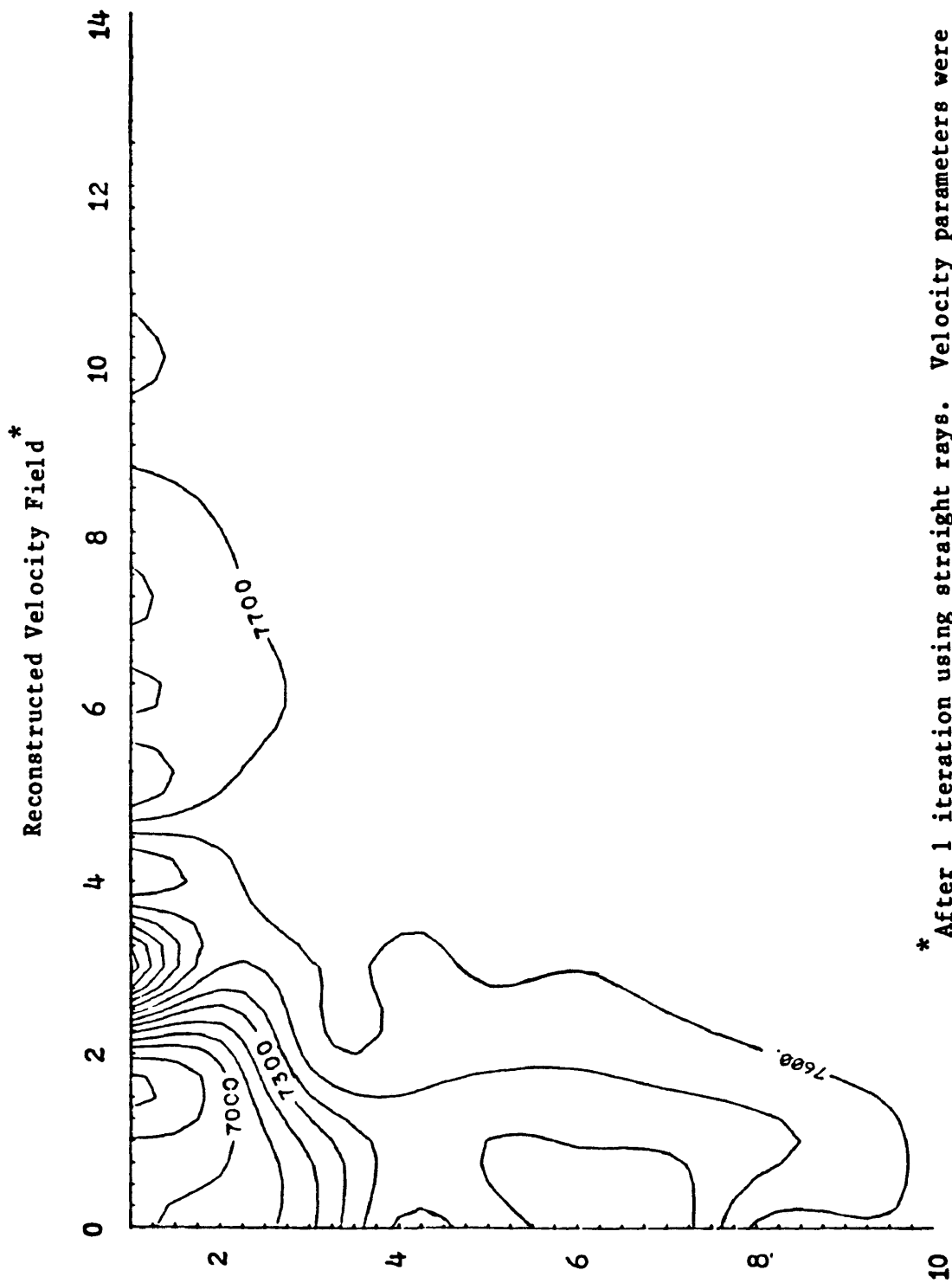
* After 1 iteration using straight rays.

Figure 17a. Acoustic tomography from the Pierre Shale VSP data.



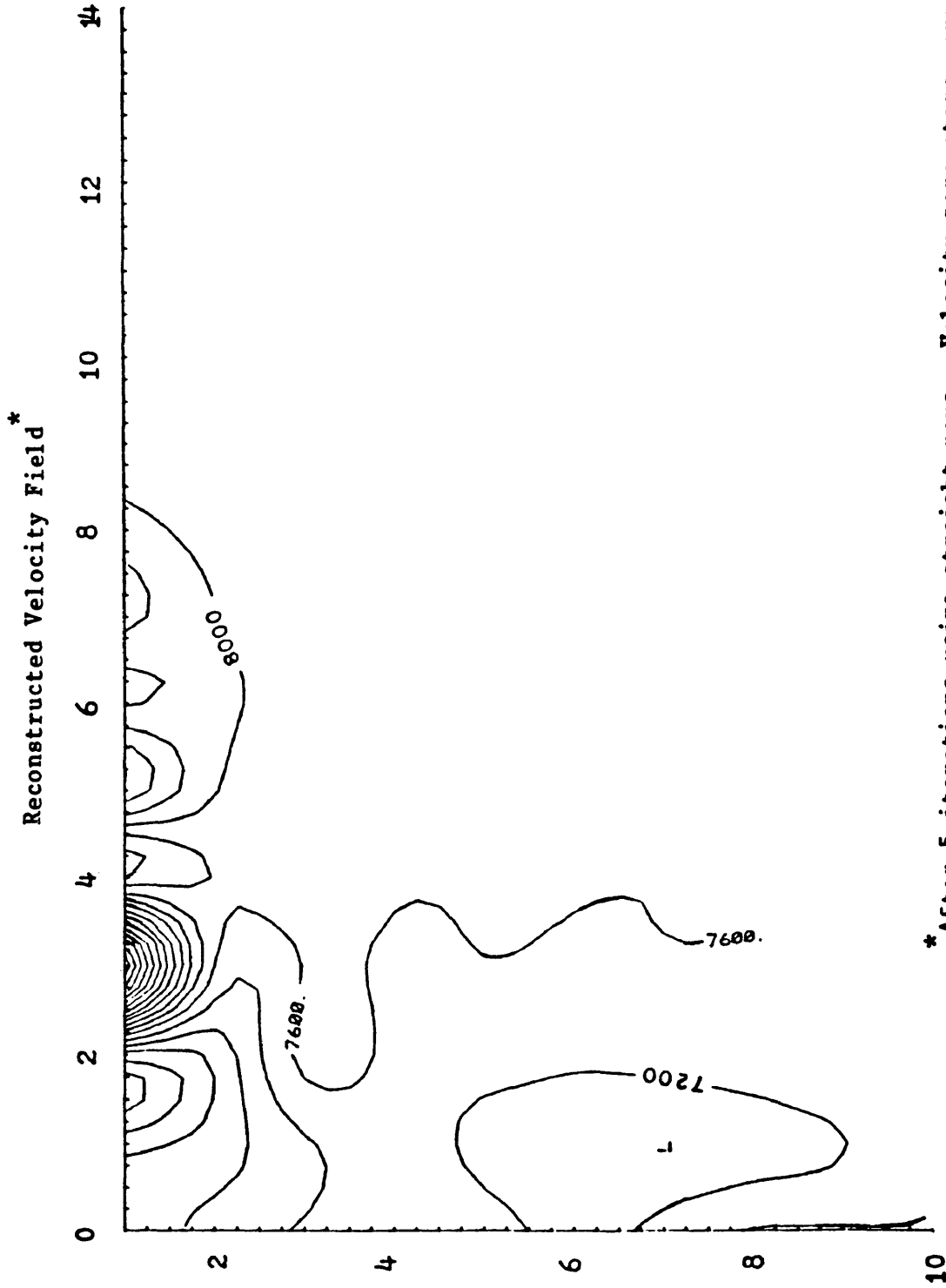
* After 5 iterations using straight rays.

Figure 17b. Acoustic Tomography from the Pierre Shale VSP data.



* After 1 iteration using straight rays. Velocity parameters were constrained down the well (z-axis).

Figure 18a. Acoustic tomography from the Pierre Shale VSP data.



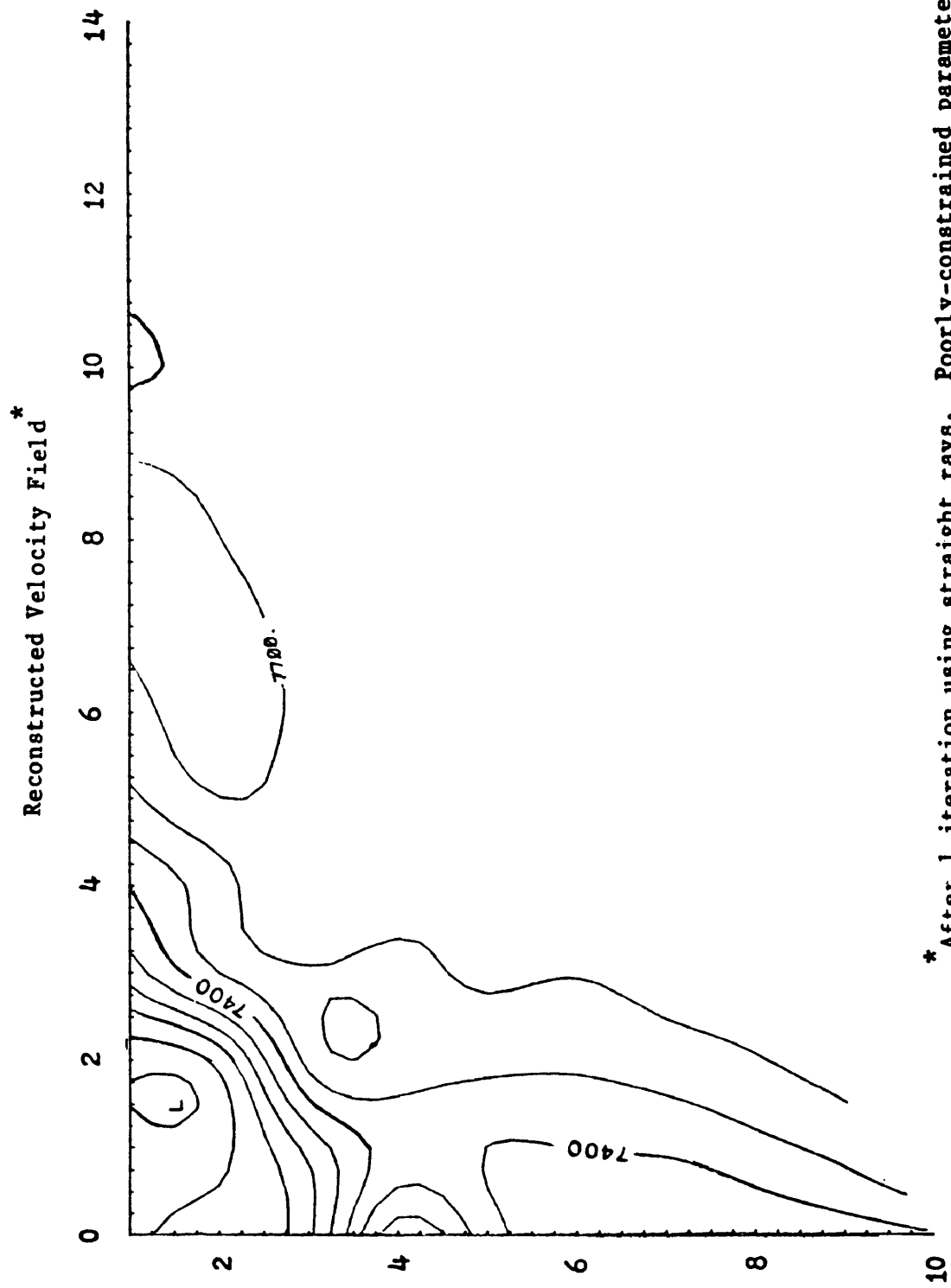
* After 5 iterations using straight rays. Velocity parameters were constrained down the well (z-axis).

Figure 18b. Acoustic tomography from the Pierre Shale VSP data.

6250 ft/sec at a depth of 100 feet, the reconstructed high, only 600 feet away, reaches 11,900 ft/sec for the unconstrained case and 12,000 ft/sec for the constrained case. However, these velocity highs consisted of a single, poorly-constrained parameter along the surface, which clearly introduces uncertainty into this velocity value.

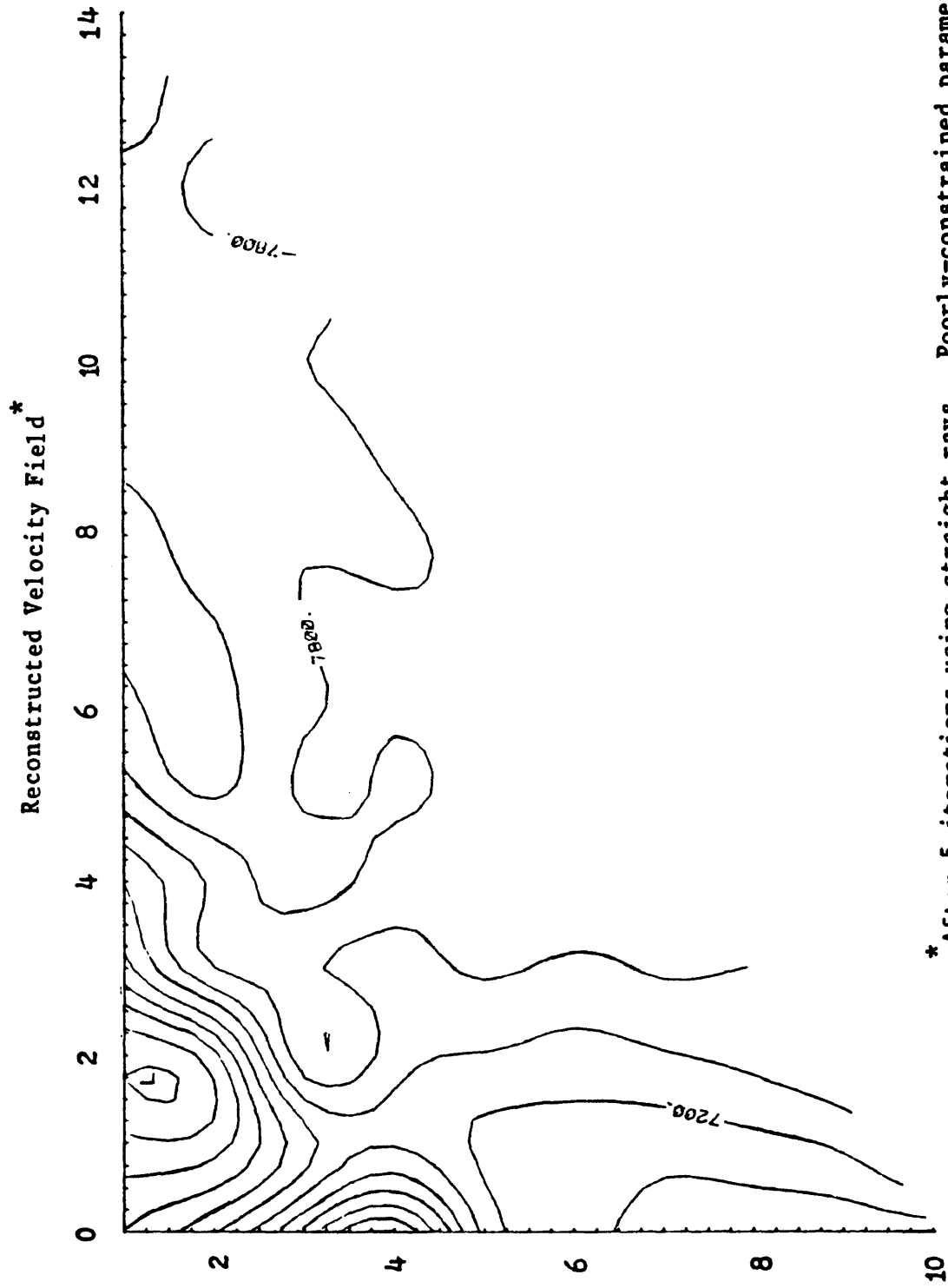
By eliminating any poorly constrained parameters along the surface, the imaging process is improved (Fig. 19, 20). These errant velocity values were ignored only when the reconstructed velocity field was contoured, and thus, they were used throughout the reconstruction algorithm. A similar step was applied to the reconstructions using circular rays based on the linear velocity trend exhibited by the sonic log (Fig. 21, 22). Each reconstruction shows a high velocity ridge of 8,000 ft/sec. hovering between a horizontal offset of 200 to 900 feet at a depth around 200 to 300 feet. Additionally, a low velocity cell exists at 100 foot offset and a depth of 500 to 800 feet. This low cell fluctuates between velocities of 7,000 to 7,200 ft/sec. depending upon the reconstruction conditions.

To test the validity of these reconstructions, the observed travel times were compared to theoretical data from a linearly increasing medium. This synthetic model, based on sonic log information, initially started at a velocity of 7,000 ft/sec. and increased to 8,000 ft/sec. at the total depth of the well. For comparison, both the synthetic and the Pierre Shale VSP travel times were plotted for each geophone posi-



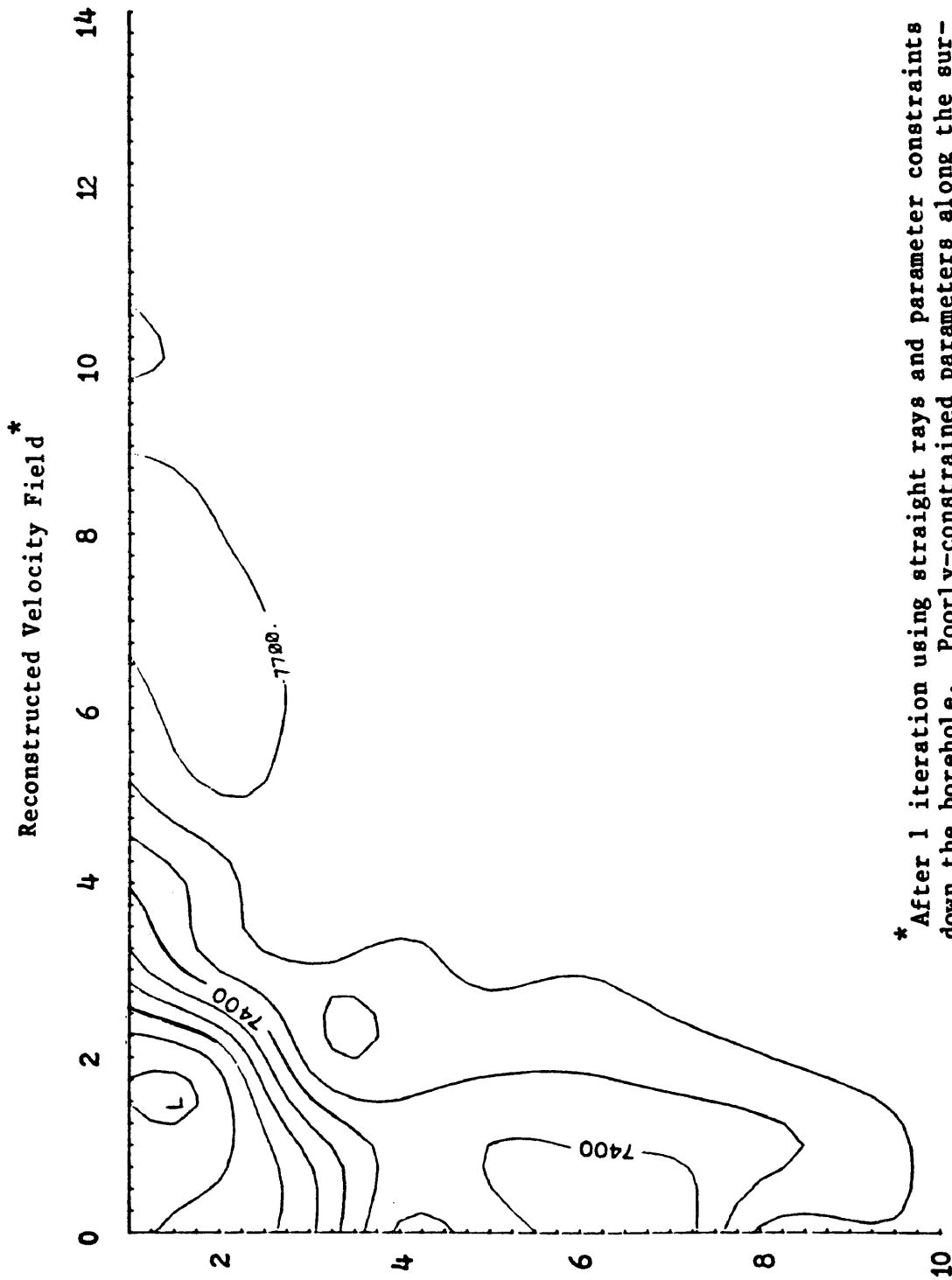
* After 1 iteration using straight rays. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 19a. Acoustic tomography from the Pierre Shale VSP data.



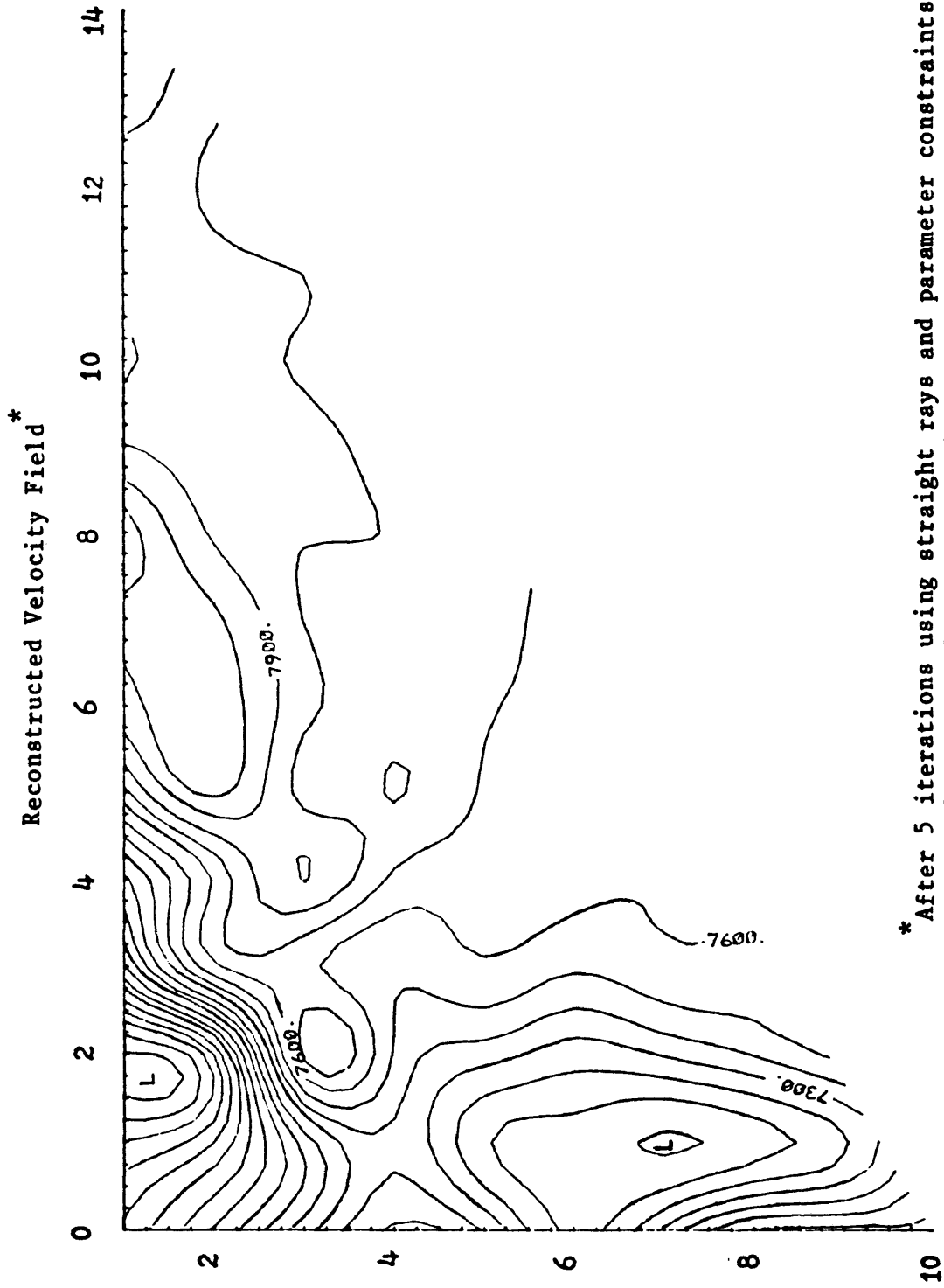
* After 5 iterations using straight rays. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 19b. Acoustic tomography from the Pierre Shale VSP data.



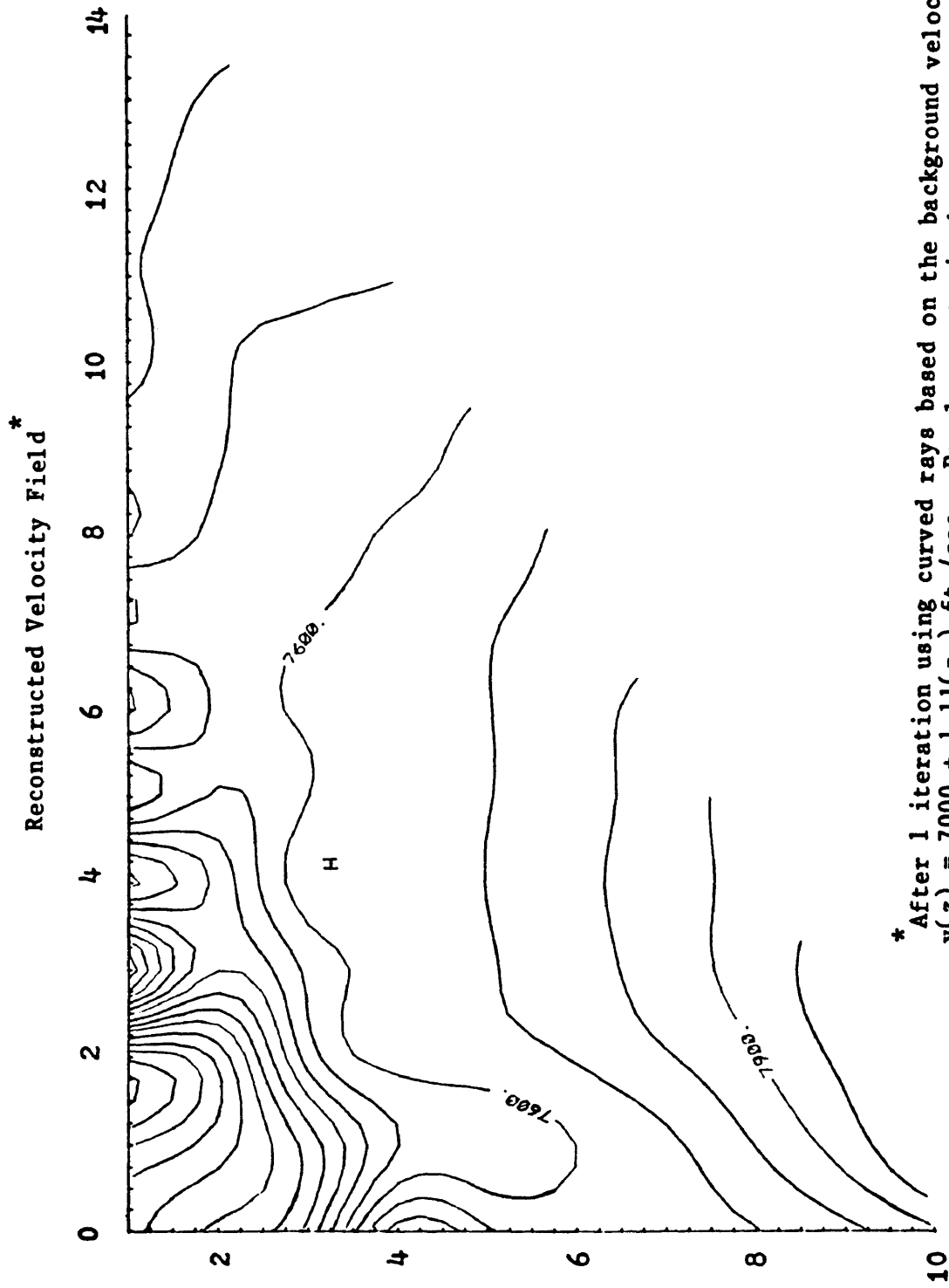
* After 1 iteration using straight rays and parameter constraints down the borehole. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 20a. Acoustic tomography from the Pierre Shale VSP data.



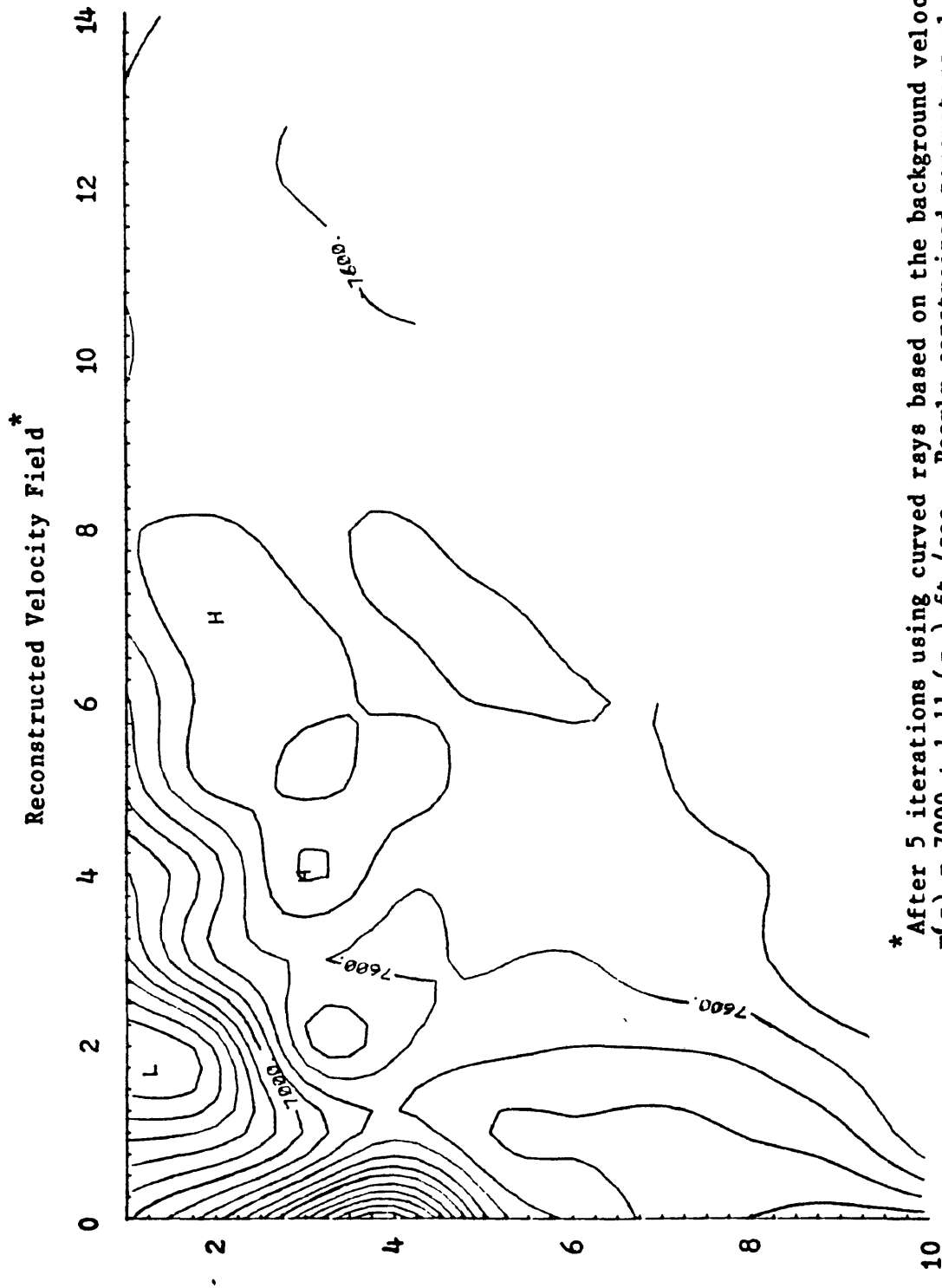
* After 5 iterations using straight rays and parameter constraints down the borehole. Poorly-constrained parameters along the surface face were eliminated from the image.

Figure 20b. Acoustic tomography from the Pierre Shale VSP data.



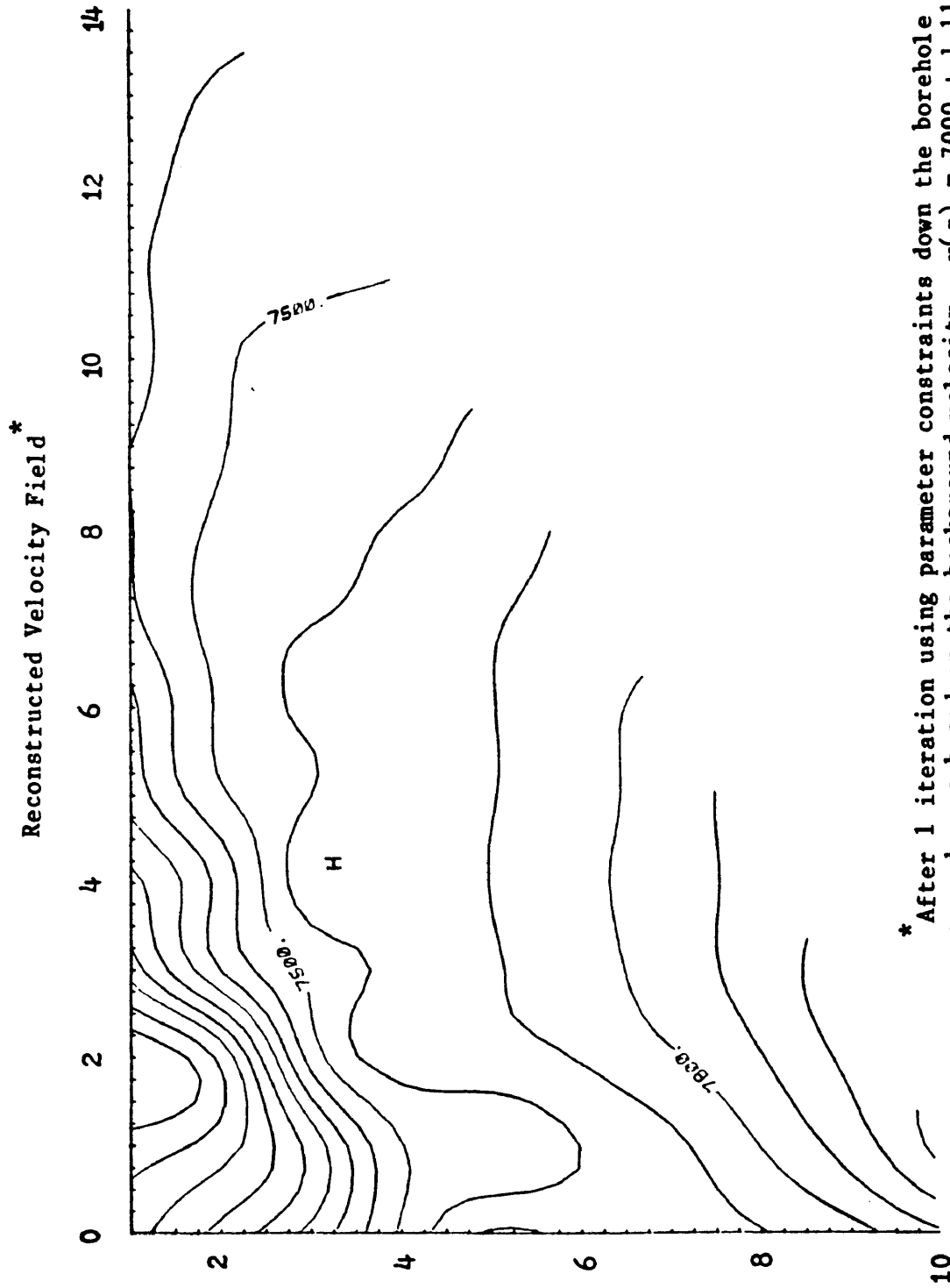
* After 1 iteration using curved rays based on the background velocity $v(z) = 7000 + 1.11(z)$ ft./sec. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 21a. Acoustic tomography from the Pierre Shale VSP data.



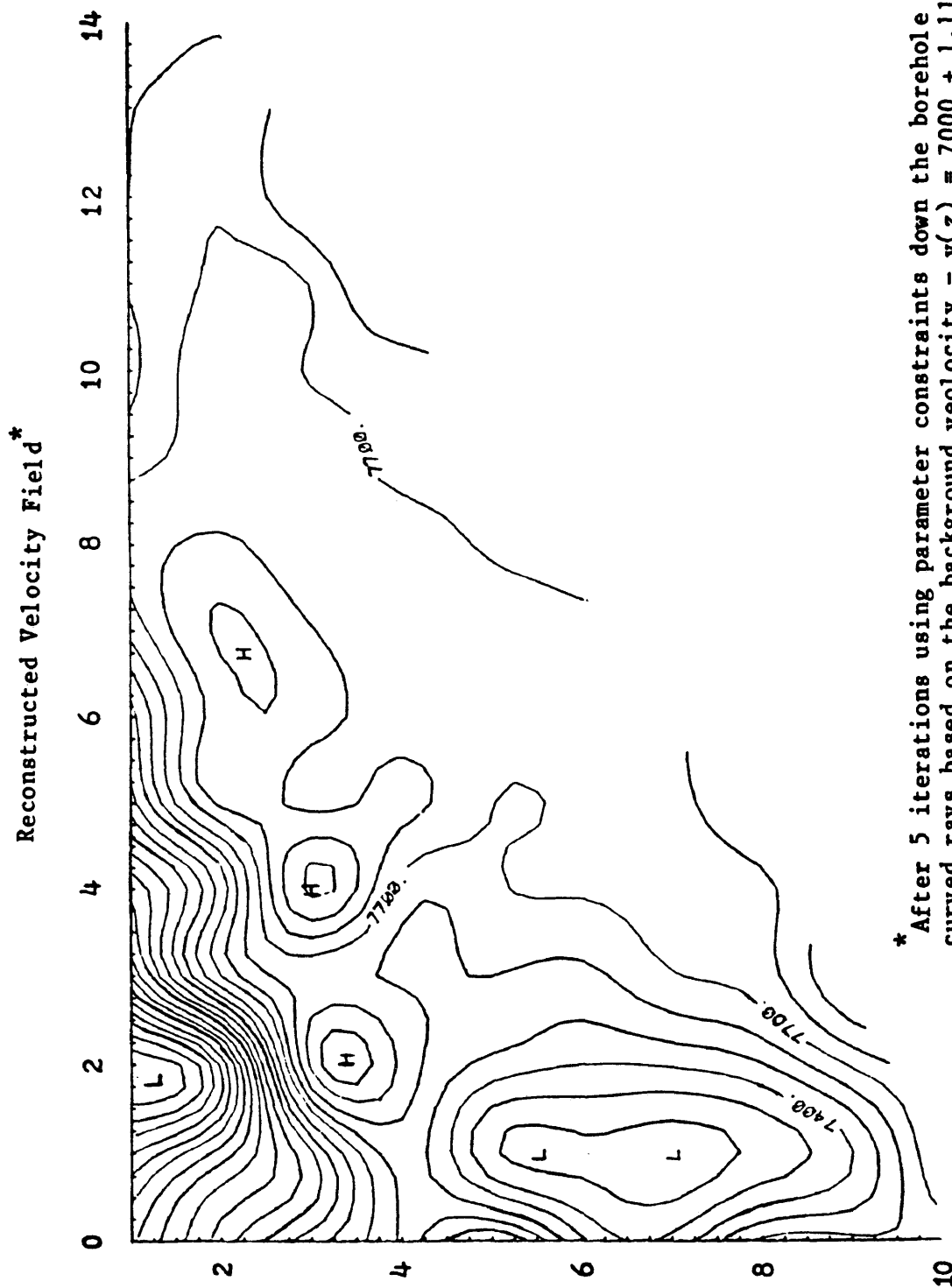
* After 5 iterations using curved rays based on the background velocity $v(z) = 7000 + 1.11(z)$ ft./sec. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 21b. Acoustic tomography from the Pierre Shale VSP data.



* After 1 iteration using parameter constraints down the borehole and curved rays based on the background velocity $-v(z) = 7000 + 1.11(z_j)$ ft./sec. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 22a. Acoustic tomography from the Pierre Shale VSP data.



* After 5 iterations using parameter constraints down the borehole and curved rays based on the background velocity $v(z) = 7000 + 1.11(z_j)$ ft./sec. Poorly-constrained parameters along the surface were eliminated from the image.

Figure 22b. Acoustic tomography from the Pierre Shale VSP data.

tion (Fig. 23). Where the synthetic curve arrives after the observed times, a velocity increase is required. Where the observed times lag behind the synthetic, lower velocities are expected. Relative to the synthetic model, we should expect higher velocities at the large offset distances, but lower velocities at small offsets and large depth.

Because surface velocities are typically very slow, the high surface velocity is contrary to expectations. Three feasible explanations arise for this velocity anomaly. Either the velocity structure exists, or the data or the collection of the data is questionable, or the basic theory is inadequate for the earth conditions exhibited by the Pierre Shale. Poorly placed shot locations could result in an artificial high at the surface, if the errant shot locations resulted in a shorter ray-path.

A more feasible explanation arises from the strong anisotropy exhibited by the Pierre Shale. After using the same data collected near Brush, Colorado, White, Martineau-Nicoletis, and Monash (1983) concluded that the horizontal velocity component was 10 to 20% faster than the vertical velocity component. Qualitatively, the rays from the large offset sources possess mostly a horizontal component and thus, should exhibit a greater than expected velocity. Similarly, the deeper traveling rays are mostly vertical and would travel at the slower velocity. Therefore, the anisotropy of the Pierre Shale explains the high surface

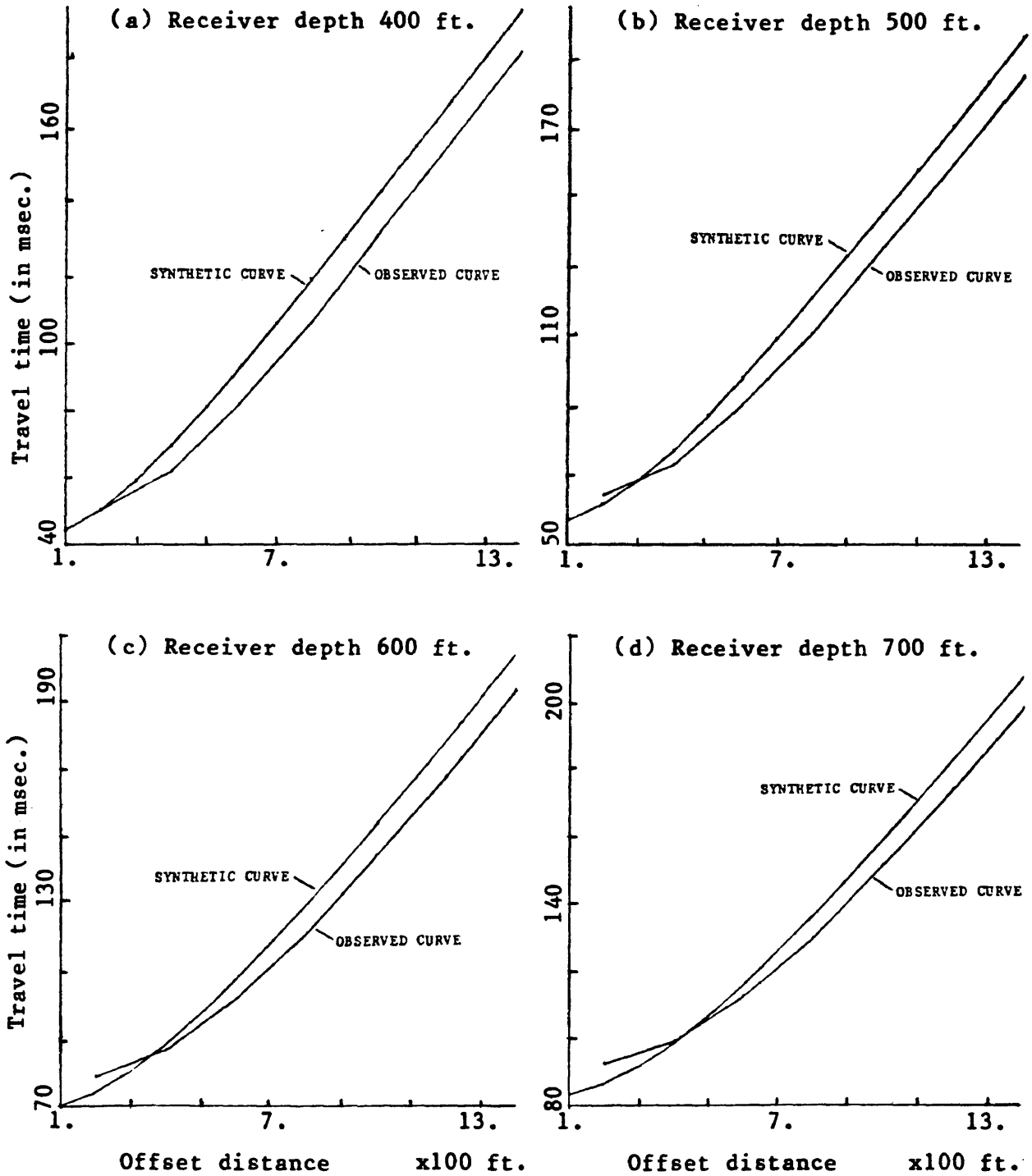


Figure 23. Travel time curves.

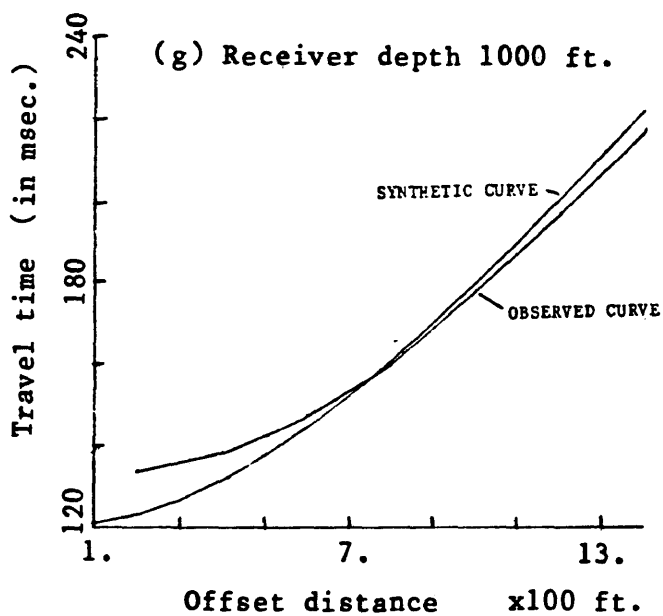
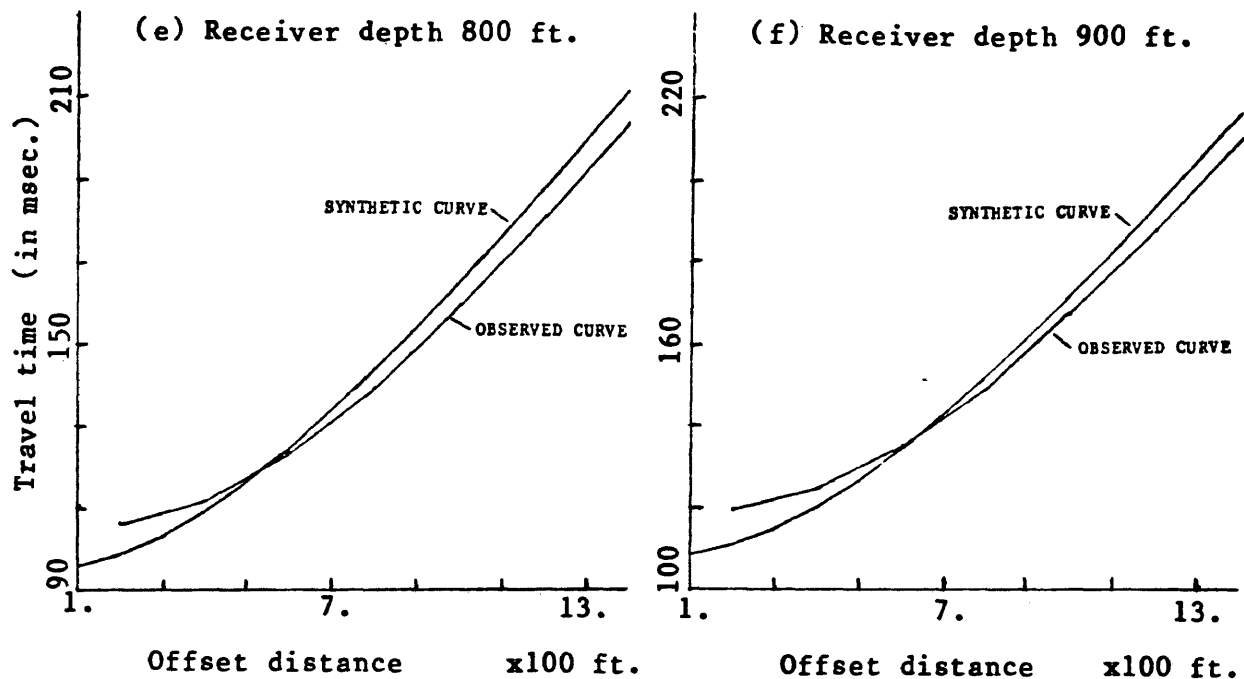


Figure 23. Travel time curves.

velocity and the low velocity cell at depth of the reconstructed velocity fields.

6. CONCLUSIONS

Algebraic reconstruction technique (ART) concentrates on the production of a reconstructed field whose projected data (travel times) agree with the observed data. This reconstructed field is modified by altering the data for each ray such that when this data is back-projected, the new image agrees with the original data.

ART possesses many inherent advantages. The flexibility of the algorithm easily allows for the incorporation of constraints. For example, sonic log information predetermines the velocity parameters down the borehole. This flexibility also permits the application of ART to any field setup of sources and receivers.

With noiseless data, ART picks a projection which agrees with the observed data. It is not in ART's nature to introduce spurious images on good data sets. However, the resolution of the image is affected by the orientation of the anomaly relative to the direction of the rays. While good resolving power occurs perpendicular to the rays, ART tends to smooth the image along the direction of the rays.

With noisy data, ART can produce single point anomalies of highly unrealistic velocities among the poorly-constrained parameters along the

edges of the model. ART will produce a reasonable reconstructed velocity field when constraints are incorporated from the well log information or by simply ignoring any poorly-constrained parameters within the image. Lastly, data, which was collected over a strongly anisotropic medium, will result in poor velocity reconstructions, because such data defies the assumption of isotropy which was used in these algorithms.

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APPENDIX A: RAY PARAMETER DERIVATION

The ray parameter within a linear $c(z)$ medium is determined from the Wiechert-Herglotz integral (Aki and Richards, 1980).

$$x_i = p \int_0^{z_j} \frac{dz}{\sqrt{\eta^2 - p^2}}, \quad (\text{A-1})$$

where $\eta = \frac{1}{c(z)}$ is the slowness,
 x_i is the source offset,
 z_j is the receiver depth.

The medium is assumed to possess a linear velocity structure.

$$c(z) = c_0 + m_0 z; \quad c_0 = c(0), \text{ the surface velocity.} \quad (\text{A-2})$$

Substituting (A-2) into (A-1), we obtain

$$x_i = p \int_0^{z_j} \frac{(c_0 + m_0 z) dz}{\sqrt{1 - p^2 (c_0 + m_0 z)^2}}. \quad (\text{A-3})$$

The integration may be carried out explicitly. The result is

$$x_i = \frac{1}{m_0 p} \sqrt{1 - p^2 c^2(z)} \Big|_{z=0}^{z=z_j}, \quad (\text{A-4})$$

$$x_i = \frac{1}{m_o p} \left[\sqrt{1 - p^2 c^2(z_j)} - \sqrt{1 - p^2 c_o^2} \right]. \quad (\text{A-5})$$

Squaring both sides, we find

$$(x_i p m_o)^2 = (1 - p^2 c_o^2) + (1 - p^2 c^2(z_j)) - 2 \sqrt{(1 - p^2 c_o^2)(1 - p^2 c^2(z_j))}, \quad (\text{A-6})$$

or

$$(x_i p m_o)^2 - 2 + p^2 (c_o^2 + c^2(z_j)) = -2 \sqrt{(1 - p^2 c_o^2)(1 - p^2 c^2(z_j))}. \quad (\text{A-7})$$

Again squaring both sides, we find

$$p^4 [m_o^2 x_i^2 + c_o^2 + c^2(z_j)]^2 - 4 c_o^2 p^2 [m_o^2 x_i^2 + c_o^2 + c^2(z_j)] + 4 = 4 (1 - p^2 c_o^2) (1 - p^2 c^2(z_j)), \quad (\text{A-8})$$

or

$$p^4 \left[[m_o^2 x_i^2 + c_o^2 + c^2(z_j)]^2 - 4 c_o^2 c^2(z_j) \right] - 4 p^2 m_o^2 x_i^2 = 0. \quad (\text{A-9})$$

Solving for p, we obtain

$$p^2 = 0, \quad p^2 = \frac{4 m_o^2 x_i^2}{[m_o^2 x_i^2 + c_o^2 + c^2(z_j)]^2 - 4 c_o^2 c^2(z_j)}. \quad (\text{A-10})$$

However, the ray parameter must be positive for a linearly increasing medium. Thus, we conclude that

$$p = \frac{2 m_0 x_i}{\sqrt{[m_0^2 x_i^2 + c_0^2 + c^2(z_j)]^2 - 4 c_0^2 c^2(z_j)}}. \quad (\text{A-11})$$

APPENDIX B: TRAVEL TIME CALCULATION

The travel time in a linear $c(z)$ medium is calculated by the Wiechert-Herglotz inversion method (Aki and Richards, 1980).

$$T_{ij} = \int_0^{z_j} \frac{dz}{c(z) \sqrt{1 - p^2 c^2(z)}}. \quad (\text{B-1})$$

We introduce the new variable of integration θ :

$$p c(z) = \sin \theta, \quad (\text{B-2})$$

$$\sqrt{1 - p^2 c^2(z)} = \cos \theta. \quad (\text{B-3})$$

Differentiating (B-3), we obtain

$$p m_0 dz = \cos \theta d\theta. \quad (\text{B-4})$$

Substituting (B-2) and (B-4) into (B-1) yields:

$$T_{ij} = \frac{1}{m_0} \int \frac{d\theta}{\sin \theta}, \quad (\text{B-5})$$

$$T_{ij} = \frac{-1}{m_0} \ln [\csc \theta + \cot \theta]. \quad (\text{B-6})$$

In terms of z , this yields the result

$$T_{ij} = \frac{1}{m_0} \ln \left[\frac{1 + \sqrt{1 - p^2 c^2(z)}}{pc(z)} \right] \Big|_{z=z_j}^{z=0}, \quad (\text{B-7})$$

$$T_{ij} = \frac{1}{m_0} \ln \left[\frac{c(z_j) (1 + \sqrt{1 - p^2 c_0^2})}{c_0 (1 + \sqrt{1 - p^2 c^2(z_j)})} \right]. \quad (\text{B-8})$$

APPENDIX C: COMPUTER PROGRAM LISTING

```

*****
*
*   VSPART: Vertical Seismic Profile inversion of travel time
*           data by Algebraic Reconstruction Technique
*
*   Programmed by KLS—May 1984
*
*   ****   VARIABLE USAGE   ****   VARIABLE USAGE   ****
*
*   C0      Initial velocity at the surface (in ft/sec)
*   C1      Increase in velocity with a 'dx' increase in depth
*           (in ft/sec-grid)
*   CMIN    Minimum acceptable chord length for ART
*   DCOM    Valid commands
*   DFILE   Data file name
*   DX      Spacing between geophones and shots (in ft)
*           also conversion factor — x ft=1 grid unit
*   ERLAST  Previous iteration's mean square error
*   ERROR   Mean square error in calculated times
*   ICOM    Inputted command
*
*   LART    .FALSE.      Use first guess ART routine
*           .TRUE.       Use iterative ART routine
*   LBACK   .FALSE.      Constant velocity background
*           .TRUE.       Linear velocity background
*
*   M       # of source locations
*   MROW    Maximum # of sources
*   N       # of receivers down borehole
*   NCOL    Maximum # of receivers
*   NDCOM   # of valid commands
*   N ITER  # of iterations performed
*   RDELTA  Length of 'dx' segment for integration
*
*   ****   COMMON STORAGE   ****
*
*   SLOW()  Slowness at parameter point m,n (in sec/grid)
*
*   TIME()  First arrival time from m th source to n th receiver
*
*   LCONST  .FALSE. Unconstrained parameters down well
*           .TRUE.  Constrained parameters down well
*   VELVAR  Allowable slowness deviation from constrained value
*   CONST() Constrained parameter values
*
*****

```

```

*****
PARAMETER      (MROW=20,NCOL=10)
IMPLICIT       LOGICAL(L)
INTEGER        L
CHARACTER      DFILE*10,      DCOM*65,      ICOM*5
COMMON/CONSTR/ LCONST,VELVAR,CONST(NCOL)
COMMON/TIMES/  TIME(MROW,NCOL)
COMMON/VELOCI/ SLOW(0:MROW,0:NCOL)
DATA DCOM/'    ART  BACKGCONSTDATA EXIT HELP KEYINMAP  PARAM
*TIME VELOC    '/
DATA  NDCOM/11/,      LART/.FALSE./,  LCONST/.FALSE./
DATA  CMIN/.1/,      RDELTA/.1/

*****  Command input  *****

TYPE 1000
10 TYPE 1001
ACCEPT '(A5)', ICOM
1000 FORMAT('1',79('*'))//T25,'Vertical Seismic Profiling'/T24,
A'Inversion of Travel Time Data'/T20,'By Algebraic Reconstruction
B Technique'//1X,79('*'))//T23,'Programmed by KLS - - July 1984')
1001 FORMAT(/1X,'5-letter command?',2X,f)

*****  Command branching  *****
*****  Accepts unique abbr. for commands  *****

CALL ABBREV      (DCOM,NDCOM,ICOM,5)
IF(ICOM.EQ.'ART  ') GOTO 200
IF(ICOM.EQ.'BACKG') GOTO 350
IF(ICOM.EQ.'CONST') GOTO 450
IF(ICOM.EQ.'DATA ') GOTO 100
IF(ICOM.EQ.'EXIT ') STOP
IF(ICOM.EQ.'HELP ') GOTO 50
IF(ICOM.EQ.'KEYIN') GOTO 150
IF(ICOM.EQ.'MAP  ') GOTO 400
IF(ICOM.EQ.'PARAM') GOTO 60
IF(ICOM.EQ.'TIME ') GOTO 250
IF(ICOM.EQ.'VELOC') GOTO 300
TYPE 1010
GOTO 10
1010 FORMAT(1X,'INVALID COMMAND - - - Type HELP for legal commands')

*****  Help command  *****

50 TYPE 1050
GOTO 10

```

```

1050  FORMAT(/1X,'***** LEGAL COMMANDS *****'/
A1X,'ART      Algebraic reconstruction technique - iterate once'/
B1X,'BACKG    Choose linear or constant velocity background'/
C1X,'CONST    Constrain the slowness parameters down the well'/
D1X,'DATA     Input travel time data from data file'/
E1X,'EXIT     End program'/
H1X,'HELP     Prints this list'/
K1X,'KEYIN    Keyin travel time data'/
M1X,'MAP      Creates a datafile for the contouring programs'/
P1X,'PARAM    Change parameters for ART'/
T1X,'TIME     Print travel time grid onto the screen'/
V1X,'VELOC    Print velocity grid onto the screen'/
Z/1X,'Accepts unique abbrev. for any commands')

```

```
***** Parameter command *****
```

```

60      TYPE 1060
        TYPE 1062μ      ACCEPT *, RDELTAμ      RDELTA=RDELTA/DX
        TYPE 1064μ      ACCEPT *, CMINμ      CMIN=SQRT(2.0)*CMIN/100.
        GOTO 10
1060    FORMAT(/1X,'Routine to change parameters used in ART')
1062    FORMAT(/1X,'In the forward model, travel times are computed by i
*ntegrating the slowness'/1X,'along the raypath.'//1X,'Enter the le
*ngth of a πdxπ segment (in feet) for integration.',2X,f)
1064    FORMAT(/1X,'In calculating the average slownesses, problems aris
*e, due to computer errors,'/1X,'whenever the raypath just barely i
*nteseacts a parameter circle. To avoid this'/1X,'problem, the data
* is ignored if the intersecting chord length is smaller than'/1X,'
*some minimum value.'//1X,'Enter this minimum value as a percentage
* of the maximum value.',2X,f)

```

```
***** Data command *****
```

```

100     TYPE 1100
        ACCEPT 1101,    DFILE
        OPEN(UNIT=1,DEVICE='DSK',FILE=DFILE,DISPOSE='SAVE')
        READ(1,1102,END=10)    M,N,DX,((TIME(I,J),I=1,M),J=1,N)
        LART=.FALSE.μ    N ITER=0
        GOTO 10
1100    FORMAT(/1X,'Input data file name (ie DATA.DAT)',2X,f)
1101    FORMAT(A10)
1102    FORMAT(T20,'ART.FOR Datafile'/'# of sources=',I2,5X,
* '# of receivers=',I2,5X,'Spacing=',F5.1,2X,'ft'/(5F10.5))

```

```
***** Keyin command *****
```

```

150     TYPE 1150μ      ACCEPT *, M
        TYPE 1151μ      ACCEPT *, N

```

```

TYPE 1152μ      ACCEPT *, DX
TYPE 1155
DO 160 I=1,M
DO 160 J=1,N
      TYPE 1160, I,J
      ACCEPT *, TIME(I,J)
160  CONTINUE
      TYPE 1100
      ACCEPT 1101,   DFILE
      OPEN(UNIT=1,DEVICE='DSK',FILE=DFILE)
      WRITE(1,1102)  M,N,DX, ((TIME(I,J),I=1,M),J=1,N)
      LART=.FALSE.μ  N ITER=0
      GOTO 10
1150  FORMAT(/1X,'KEYIN DATA ROUTINE'//1X,'Input # of sources ',f)
1151  FORMAT(/1X,'Input # of receivers',2X,f)
1152  FORMAT(/1X,'Input spacing (ft) between geophones/shots',2X,f)
1155  FORMAT(/1X,'Input first arrival times - Enter zero (0.0) for
      *no data')
1160  FORMAT(1X,'Source=',I2,5X,'Receiver=',I2,5X,'Time=?',2X,f)

*****  ART command  *****

200  IF(LBACK)      THEN
      IF(LART)      THEN
          CALL ARTLB2(M,N,RDELTA,CMIN,CO/DX,C1/DX,ERROR)
      ELSE
          CALL ARTLB1(M,N,RDELTA,CMIN,CO/DX,C1/DX,ERROR)
          LART=.TRUE.
      END IF
  ELSE
      IF(LART)      THEN
          CALL ART2      (M,N,RDELTA,CMIN,ERROR)
      ELSE
          CALL ART1      (M,N,RDELTA,CMIN,ERROR)
          LART=.TRUE.
      END IF
  END IF
  TYPE 1200,      ERROR
  N ITER=N ITER + 1
  IF(N ITER.GT.1) TYPE 1210,      ERLAST
  ERLAST=ERROR
  GOTO 10
1200  FORMAT(/1X,'Mean square error is ',F12.7)
1210  FORMAT(1X,'Last error value was ',F12.7)

*****  Time command  *****

250  TYPE 1250

```

```

      IF(M.LE.9)      THEN
          M1=1μ      M2=M
      ELSE
          TYPE 1301μ      ACCEPT *, M1
          TYPE 1302μ      ACCEPT *, M2
      END IF
      IF(N.LE.20)     THEN
          N1=1μ      N2=N
      ELSE
          TYPE 1303μ      ACCEPT *, N1
          TYPE 1304μ      ACCEPT *, N2
      END IF
      TYPE 1310, (I,I=M1,M2)μ      TYPE 1311
      DO 260 J=N1,N2
          TYPE 1260,      J, (TIME(I,J),I=M1,M2)
260    CONTINUE
      GOTO 10
1250   FORMAT(/1X,'TRAVEL TIME INPUT'/)
1260   FORMAT(1X,I2,3X,14F8.5)

***** Velocity command *****

300    TYPE 1300
      IF(M.LE.8)      THEN
          M1=0μ      M2=M
      ELSE
          TYPE 1301μ      ACCEPT *, M1
          TYPE 1302μ      ACCEPT *, M2
      END IF
      IF(N.LE.19) THEN
          N1=0μ      N2=N
      ELSE
          TYPE 1303μ      ACCEPT *, N1
          TYPE 1304μ      ACCEPT *, N2
      END IF
      TYPE 1305, N ITER
      TYPE 1310, (I,I=M1,M2)μ      TYPE 1311
      DO 310 J=N1,N2
          TYPE 1312,      J, (VEL(DX,SLOW(I,J)),I=M1,M2)
310    CONTINUE
      GOTO 10
1300   FORMAT(/1X,'***** Routine to Print Velocity Structure *****')
1301   FORMAT(/1X,'Input starting shot point to print',2X,f)
1302   FORMAT(/1X,'Input ending shot point to print',2X,f)
1303   FORMAT(/1X,'Input starting geophone to print',2X,f)
1304   FORMAT(/1X,'Input ending geophone to print',2X,f)
1305   FORMAT(/T5,'***** Velocity Structure After',I3,' Iterations',
          *3X,5(' '))

```

```

1310  FORMAT(/T20,'Shot Points'/T11,14(I2,6X))
1311  FORMAT(1X,'Depth')
1312  FORMAT(1X,I2,3X,14F8.0)

*****  Background command      *****

350   TYPE 1350
      ACCEPT '(A5)', ICOM $\mu$     CALL UCASE2(.TRUE.,5,ICOM)
      IF(ICOM(1:1).EQ.'C')    GOTO 360
      IF(ICOM(1:1).EQ.'L')    GOTO 370
      GOTO 350

360   LBACK=.FALSE. $\mu$   GOTO 10
370   LBACK=.TRUE.
      TYPE 1370 $\mu$       ACCEPT *, C0
      TYPE 1380 $\mu$       ACCEPT *, C1
      TYPE 1390
      GOTO 10

1350  FORMAT(/1X,'CONSTANT or LINEAR velocity background?',2X,f)
1370  FORMAT(/1X,'Input initial velocity (in feet/sec)',2X,f)
1380  FORMAT(/1X,'Input velocity increase w/ geophone spacing',2X,f)
1390  FORMAT(/T20,'*** WARNING ***'/1X,'If the velocity background
      *is nearly linear, then variable overflows may occur'//
      *T20,'*** WARNING ***'/1X,'If the linear velocity background is
      *too steep, then the rays will turn and'/1X,'travel upwards.
      *Square roots of negative numbers occur.')
```

***** Map command *****

```

400   TYPE 1400
      ACCEPT '(A10)', DFILE
      OPEN(UNIT=1,DEVICE='DSK',FILE=DFILE)
      DO 410 K=0,M
      DO 410 L=0,N
      IF(SLOW(K,L).LE.0.0)    GOTO 410
      WRITE(1,*)             FLOAT(K),          -FLOAT(L),          DX/SLOW(K,L)
410   CONTINUE
      CLOSE(UNIT=1)
      GOTO 10

1400  FORMAT(/1X,'***** MAPPING ROUTINE *****'//1X,
      *'A datafile is created for input into the contouring programs'
      *//T10,'MNL:KGRID, MNL:FIT, and MNL:TOPO'
      *//1X,'Enter datafile name.ext (ie KGRID.DAT)',2X,f)
```

***** Constrain parameters command *****

```

450   TYPE 1450
      ACCEPT '(A5)', ICOM $\mu$     CALL UCASE2(.TRUE.,5,ICOM)
      IF(ICOM(1:1).EQ.'U')    GOTO 460
```

```

      IF(ICOM(1:1).EQ.'C')      GOTO 470
      GOTO 450
460   LCONST=.FALSE.μ GOTO 10
470   LCONST=.TRUE.
      TYPE 1470
      DO 480 J=1,N
          TYPE 1480, J*DX
          ACCEPT *, S
          CONST(J)=DX*S/1000000.0 ! in sec/grid units
480   CONTINUE
      TYPE 1490
      ACCEPT *, VELVAR
      GOTO 10
1450  FORMAT(/1X,'CONStrain or UNCONstrain the slowness parameters
* down the well?',2X,f)
1470  FORMAT(/1X,'Input the slowness in micro seconds per foot for
* the depth given'/1X,'Enter zero for an unconstrained parameter')
1480  FORMAT(T10,'depth',F6.0,1X,'ft',5X,'slowness?',2X,f)
1490  FORMAT(/1X,'Enter the allowable deviation of these constrained
* parameters'/1X,'(ie 0.0 for a completely constrained parameters'/
*/1X,'Enter value in percent',2X,f)

      END

```

SUBROUTINE ART1 (MROW,NCOL,DELTAR,CMIN,ERROR)

```

*****
*
*   ART: Algebraic Reconstruction Technique
*           Programmed by KLS  March 1984
*
*   REF: Mason, I. M., πAlgebraic Reconstruction of a Two
*         Dimensional Velocity Inhomogeneity in the High Hazles
*         Seam of Thoresby Collieryπ Geophysics 46 (1981) 298-308*
*
*   ***** CALLING VARIABLES *****
*
*   MROW  # of sources along the x axis
*   NCOL  # of receivers along the z axis
*   DELTAR  Approx length of ray segment for integration
*   CMIN   Minimum acceptable chord length to use to calculate
*           the average slownesses
*
*   ***** OUTPUT VARIABLES *****
*
*   ERROR  Mean square error
*
*****

```

```

*          ***** COMMON STORAGE ***** INPUT *****          *
*
*          LCONST Logical switch for constraining the slowness parameters*
*                  down the borehole (known from the sonic log)      *
*
*          TIME() First arrival times from m'th source to n'th geophone *
*                  Zero travel time signifies a lack of data          *
*
*          ***** COMMON STORAGE ***** OUTPUT *****          *
*
*          SLOW(M,N)      Slowness array                               *
*
*          TCALC(M,N)     Calculated travel times                     *
*
*-----*
*          PARAMETER      (M=20,N=10)
*          LOGICAL        LCONST
*          DIMENSION      CSUM(0:M,0:N),TSUM(0:M,0:N)
*          COMMON/CONSTR/  LCONST
*          COMMON/VELOCI/  SLOW(0:M,0:N)
*          COMMON/CALCUL/  TCALC(0:M,0:N)
*          COMMON/TIMES/   TIME(M,N)
*          EQUIVALENCE    (TSUM(0,0),TCALC(0,0))
*-----*
*          ***** VARIABLE USAGE ***** VARIABLE USAGE *****   *
*
*          A      X-squared coefficient of quadratic equation        *
*          B      X coefficient of quadratic equation                *
*          C      Coefficient of quadratic equation                  *
*          CSUM() Sum of chord lengths around m,n parameter          *
*          CLEN   Chord length from X1,Z1 to X2,Z2                  *
*          ERR    Residual error                                     *
*          I      Position of source                                 *
*          J      Position of reciever                               *
*          K      X parameter coordinate                            *
*          L      Z parameter coordinate                            *
*          NSEG   # of ray segments                                  *
*          RAD    Radical of quadratic solution                    *
*          RLEN   Length of ray path from ith source to jth receiver *
*          SEG    True length of ray path                           *
*          TSUM() Sum of times along chord lengths around m,n        *
*          X,Z    Coordinates of midpoint of ray segment            *
*          X1,Z1  Point of chord on circle                          *
*          X2,Z2  Second point of chord on circle                   *
*
*          ***** UNITS OF DIMENSIONS *****                       *
*

```

```

*      Distance      in grid units (spacing between geophones)  *
*      Time          in seconds                                     *
*      Slowness      in seconds/grid unit                         *
*
*****
*****  Clear sumations *****

DO 05  K=0,MROW
DO 05  L=0,NCOL
        TSUM(K,L)=0.0
        CSUM(K,L)=0.0
05     CONTINUE

*****  For each ray path... *****

DO 20  I=1,MROW
DO 20  J=1,NCOL

*****  For no data on ray *****

        IF(TIME(I,J).LE.0.0)  GOTO 20

*****  Ray length *****

        RLEN=SQRT(FLOAT(I*I + J*J))

*****  For each parameter circle... *****

        ZRANGE=1.0 / (SQRT(2.0)*SIN(ATAN(I/FLOAT(J))))
DO 10  K=0,I
        Z=J - J*K/FLOAT(I)
        L1=NINT(Z-ZRANGE+0.5)μ  IF(L1.LT.0)  L1=0
        L2=Z+ZRANGEμ  IF(L2.GT.J)  L2=J
DO 10  L=L1,L2

*****  Find intersection of two equations *****
*****  Equation one: Circle (X-K)**2 + (Z-L)**2 = 1/2
*****  Equation two: Line  Z = J - J*X/I
*****  Solve quadratic equation  AX**2 + 2BX + C

        A=1.0 + J*J/FLOAT(I*I)
        B=(L-J)*J/FLOAT(I) - K
        C=K*K + (J-L)**2 - 0.5
        RAD=B*B - A*C

*****  Imaginary solution means the lines do not intersect *****

```

```

                IF(RAD.LE.0.0) GOTO 10
                RAD=SQRT(RAD)

***** Locate intersection points *****
***** B.C. - Ray path starts at (I,0) and ends at (0,J) *****

                IF(K.EQ.I .AND. L.EQ.0) THEN
                    X1=K $\mu$    Z1=0
                ELSE
                    X1=(-B+RAD)/A
                    Z1=J-J*X1/I
                END IF
                IF(K.EQ.0 .AND. L.EQ.J) THEN
                    X2=0 $\mu$    Z2=L
                ELSE
                    X2=(-B-RAD)/A
                    Z2=J-J*X2/I
                END IF

***** Chord length *****

                CLEN=SQRT((X1-X2)**2 + (Z1-Z2)**2)
                IF(CLEN.LT.CMIN) GOTO 10

***** Summation of parameter times *****

                TSUM(K,L)=TSUM(K,L)+CLEN*TIME(I,J)/RLEN

***** Summation of chord lengths *****

                CSUM(K,L)=CSUM(K,L)+CLEN
10          CONTINUE
20          CONTINUE

***** Average slowness at all points *****

                DO 30 K=0,MROW
                DO 30 L=0,NCOL
                    IF(CSUM(K,L).EQ.0.0) THEN
                        SLOW(K,L)=0.0
                    ELSE
                        SLOW(K,L)=TSUM(K,L)/CSUM(K,L)
                    END IF
30          CONTINUE

***** Constrained parameters *****

                IF(LCONST) CALL PARCON(NCOL)

```

```

***** For each ray path... *****
DO 50 I=1,MROW
DO 50 J=1,NCOL
      IF(TIME(I,J).LE.0.0) GOTO 50
      TCALC(I,J)=0.0

***** Divide ray into equal segments *****

      RLEN=SQRT(FLOAT(I*I + J*J))
      NSEG=INT(RLEN/DELTAR)
      SEG=RLEN/NSEG

***** Interpolate slowness at midpoint of segments *****

      XINC=FLOAT(I)/FLOAT(NSEG)
      XSTART=XINC/2.0
      XEND=FLOAT(I)
      DO 40 X=XSTART,XEND,XINC
          Z=J-J*X/FLOAT(I)
          TCALC(I,J)=TCALC(I,J) + CARINT(X,Z)*SEG
40      CONTINUE
50      CONTINUE

***** Mean square error *****

      ERROR=0.0
      DO 60 I=1,MROW
      DO 60 J=1,NCOL
          IF(TIME(I,J).LE.0.0) GOTO 60
          ERR=TCALC(I,J)-TIME(I,J)
          ERROR=ERROR+ERR*ERR
60      CONTINUE
      RETURN
      END

```

SUBROUTINE ART2 (MROW,NCOL,DELTAR,CMIN,ERROR)

```

*****
*
*   ART2: Algebraic Reconstruction Technique
*           iterative portion
*           Programmed by KLS March 1984
*
*   REF: Mason, Geophysics 1981
*
*   ***** CALLING VARIABLES *****
*

```

```

*
*      MROW      # of sources along the x axis
*      NCOL      # of receivers along the z axis
*      DELTAR    Approx. length of ray segment for integration
*      CMIN      Minimum acceptable chord length to use to calculate
*                average slownesses
*
*      *****  OUTPUT VARIABLES      *****
*
*      ERROR     Mean square error
*
*      *****  COMMON STORAGE *****  INPUT *****
*
*      LCONST    Logical switch for constraining the slowness parameters
*                down the borehole (known from the sonic log)
*
*      TIME()    First arrival times from m'th source to n'th receiver
*                Zero travel time signifies a lack of data
*
*      *****  COMMON STORAGE *****  I/O *****
*
*      SLOW(M,N)      Slowness array
*
*      TCALC(M,N)     Calculated travel times
*

```

```

*****
PARAMETER      (M=20,N=10)
LOGICAL        LCONST
DIMENSION      CSUM(0:M,0:N),ESUM(0:M,0:N)
COMMON/CONSTR/ LCONST
COMMON/VELOCI/ SLOW(0:M,0:N)
COMMON/CALCUL/ TCALC(0:M,0:N)
COMMON/TIMES/  TIME(M,N)
*****

```

```

*
*      *****  VARIABLE USAGE *****  VARIABLE USAGE *****
*
*      A      X-squared coefficient of quadratic equation
*      B      X coefficient of quadratic equation
*      C      Coefficient of quadratic equation
*      CSUM() Sum of chord lengths around m,n parameter
*      CLEN    Chord length from x1,z1 to x2,z2
*      ERR     Residual error
*      ESUM()  Sum of errors associated with parameter m,n
*      I      Position of source
*      J      Position of reciever
*      K      X parameter coordinate
*      L      Z parameter coordinate
*

```

```

*      NSEG      # of ray segments
*      RAD       Radical of quadratic solution
*      RLEN      Length of ray path from ith source to jth receiver
*      SEG       True length of ray path
*      X,Z       Coordinates of midpoint of ray segment
*      X1,Z1     Point of chord on circle
*      X2,Z2     Second point of chord on circle
*
*      ***** UNITS OF DIMENSIONS *****
*
*      Distance      in grid units (spacing between geophones)
*      Time          in seconds
*      Slowness      in seconds/grid unit
*
*****
***** Clear summations *****
      DO 10  K=0,MROW
      DO 10  L=0,NCOL
             CSUM(K,L)=0.0
             ESUM(K,L)=0.0
10      CONTINUE

***** For each ray path... *****
      DO 30  I=1,MROW
      DO 30  J=1,NCOL

***** For no data on ray *****
             IF(TIME(I,J).LE.0.0) GOTO 30
             ERR=TIME(I,J)-TCALC(I,J)

***** Ray length *****
             RLEN=SQRT(FLOAT(I*I + J*J))

***** For each parameter... *****
             ZRANGE=1.0 / (SQRT(2.0)*SIN(ATAN(I/FLOAT(J))))
      DO 20  K=0,I
             Z=J-J*K/FLOAT(I)
             L1=NINT(Z-ZRANGE+0.5)μ IF(L1.LT.0) L1=0
             L2=Z+ZRANGEμ IF(L2.GT.J) L2=J
      DO 20  L=L1,L2

*      Find intersection of two equations

```

* Equation one: Circle $(X-K)**2 + (Z-L)**2 = 1/2$
 * Equation two: Line $Z = J - J*X/I$
 * Solve Quadratic Equation $AX**2 + 2BX + C$

A=1.0 + J*J/FLOAT(I*I)
 B=(L-J)*J/FLOAT(I) - K
 C=K*K + (J-L)**2 - 0.5
 RAD=B*B - A*C

***** Imaginary solution means the lines do not intersect *****

IF(RAD.LE.0.0) GOTO 20
 RAD=SQRT(RAD)

***** Locate intesection points *****
 ***** B.C. - Ray path starts at (I,0) and ends at (0,J) *****

IF(K.EQ.I .AND. L.EQ.0) THEN
 X1=K_μ Z1=0
 ELSE
 X1=(-B+RAD)/A
 Z1=J-J*X1/I

END IF
 IF(K.EQ.0 .AND. L.EQ.J) THEN
 X2=0_μ Z2=L
 ELSE
 X2=(-B-RAD)/A
 Z2=J-J*X2/I
 END IF

***** Chord length *****

CLEN=SQRT((X1-X2)**2 + (Z1-Z2)**2)
 IF(CLEN.LT.CMIN) GOTO 20

***** Summation of chord lengths *****

CSUM(K,L)=CSUM(K,L) + CLEN

***** Summation of error around parameter *****

ESUM(K,L)=ESUM(K,L) + ERR*CLEN/RLEN

20 CONTINUE

30 CONTINUE

***** Update parameter matrix *****

DO 40 K=0,MROW

```

DO 40  L=0,NCOL
      IF(CSUM(K,L).EQ.0.0) THEN
          SLOW(K,L)=0.0
      ELSE
          SLOW(K,L)=SLOW(K,L) + ESUM(K,L)/CSUM(K,L)
      END IF
40  CONTINUE

***** Constrained parameters *****

IF(LCONST) CALL PARCON(NCOL)

***** For each ray path... *****

DO 60  I=1,MROW
DO 60  J=1,NCOL
      IF(TIME(I,J).LE.0.0) GOTO 60
      TCALC(I,J)=0.0

***** Divide ray into equal segments *****

      RLEN=SQRT(FLOAT(I*I + J*J))
      NSEG=INT(RLEN/DELTAR)
      SEG=RLEN/NSEG

***** Interpolate slowness at midpoint of segment *****

      XINC=FLOAT(I)/FLOAT(NSEG)
      XSTART=XINC/2.0
      XEND=FLOAT(I)
      DO 50  X=XSTART,XEND,XINC
          Z=J-J*X/FLOAT(I)
          TCALC(I,J)=TCALC(I,J) + CARINT(X,Z)*SEG
50  CONTINUE
60  CONTINUE

***** Mean square error *****

ERROR=0.0
DO 70  I=1,MROW
DO 70  J=1,NCOL
      IF(TIME(I,J).LE.0.0) GOTO 70
      ERR=TCALC(I,J)-TIME(I,J)
      ERROR=ERROR+ERR*ERR
70  CONTINUE
END

```

```

SUBROUTINE ARTLB1      (MROW,NCOL,DELTAR,CMIN,CO,C1,ERROR)
*****
*
*   ARTLB1: Algebraic Reconatrusion Technique for a linear
*           velocity background (ie. curved raypaths)
*
*   Programmed by KLS—May 1984
*
*   ***** CALLING VARIABLES *****
*
*   MROW      # of sources along x-axis
*   NCOL      # of recievers along z-axis
*   DELTAR    Approx. length of ray segment for intergration
*   CMIN      Minimum acceptable chord length to use to calculate
*             average slownesses
*   CO        Velocity at surface
*   C1        Change in velocity with a 'dx' increase in depth
*
*   ***** OUTPUT VARIABLES *****
*
*   ERROR     Mean square error of residual travel times
*
*   ***** COMMON STORAGE ***** INPUT *****
*
*   LCONST    Logical switch for constraining the slowness parameters*
*             down the borehole (known from the sonic log)
*
*   TIME()    First arrival times from the mth shot to the nth
*             geophone
*             Zero travel time signifies a lack of data
*
*   ***** COMMON STORAGE ***** OUTPUT *****
*
*   SLOW()    Slowness array
*
*   TCALC()   Calculated travel times
*
*****
PARAMETER      (M=20,N=10)
LOGICAL        LCONST
DIMENSION      CSUM(0:M,0:N),TSUM(0:M,0:N)
COMMON/CONSTR/ LCONST
COMMON/VELOCI/ SLOW(0:M,0:N)
COMMON/TIMES/  TIME(M,N)
COMMON/CALCUL/ TCALC(0:M,0:N)
EQUIVALENCE   (TCALC(0,0),TSUM(0,0))
*****

```

```

*          ***** VARIABLE USAGE ***** VARIABLE USAGE *****
*
* Alpha   Angle of ray path
* Angle i Angle of incidence for ray at K,0
* Beta    Angle of intersection
* CSUM()  Sum of chord lengths around m,n parameter
* CLEN    Chord (arc) length intersecting parameter circle
* DIST    Distance between parameter point and origin of
*          the radius of curvature for the raypath
*
* ERR     Residual error
* I       Position of source
* J       Position of reciever
* K       X parameter coordinate
* L       Z parameter coordinate
* NSEG    # of ray segments
* R       The radius of curvature for the ray
* RAYPAR  Ray parameter for the ray from ith shot to jth geophone
* RLEN    Length of ray path from i th source to j th receiver
* SEG     True length of ray path
* TSUM(M,N) Sum of times along chord lengths around M,N
* X,Z     Midpoint of segment for integration
* Xo,Zo   Origin of the radius of curvature
*          always a negative value
*
*          ***** UNITS OF DIMENSIONS *****
*
* Angles      in radians
* Distance    in grid units (spacing between geophones)
* Time        in seconds
* Slowness    in seconds/grid unit
* Velocity    in grid units/second
*
*****

```

```

***** Clear summations *****

```

```

DO 10 K=0,MROW
DO 10 L=0,NCOL
      CSUM(K,L)=0.0
      TSUM(K,L)=0.0
10 CONTINUE

```

```

***** For each ray path... *****

```

```

DO 30 I=1,MROW
DO 30 J=1,NCOL

```

```

***** For no data on ray *****

```

```

                IF(TIME(I,J).LE.0.0)      GOTO 30

***** Find radius of curvature for the ray *****

RAYPAR=2.0*C1*I / SQRT((C0*C0+(C0+J*C1)**2+C1*C1*I*I) **2 - 4*C0
* *C0*(C0+C1*J)**2 )
R=1.0/(RAYPAR*C1)

***** Find estimated travel time *****

T est=LOG((C0+C1*J)*(1+ SQRT(1 - (C0*RAYPAR)**2)) /
*      (C0 + SQRT(C0*C0 - (C0*(C0+C1*J)*RAYPAR)**2))) / C1

***** Use geometry *****

alpha=2.0*ASIN(0.5*SQRT(FLOAT(I*I + J*J))/R)
angle i=      ATAN(I/FLOAT(J)) - alpha/2.0
RLEN=R*Alpha
Xo=-R*COS(angle i + alpha)
Zo=-R*SIN(angle i)

***** For each parameter circle... *****

DO 20  K=0,I
      L1=NINT(Zo + SQRT((R-1./SQRT(2.))**2 - (K-Xo)**2) + .5)
      IF(L1.LT.0)      L1=0
      L2=INT(Zo + SQRT((R+1./SQRT(2.))**2 - (K-Xo)**2))
      IF(L2.GT.J)      L2=J
DO 20  L=L1,L2

* Find intersection of two equations
* Equation one: Circle (X-K)**2 + (Z-L)**2 = 1/2
* Equation two: Circle (X-Xo)**2 + (Z-Zo)**2 = R**2

***** Ray intersects parameter circle ??? *****

      DIST=SQRT((K-Xo)**2+(L-Zo)**2)
      IF(ABS(DIST - R) .GE. 1./SQRT(2.0))      GOTO 20

***** Locate intersection angle *****

***** Law of cosines *****

***** TEMP avoids computer roundoff errors *****

      TEMP=(R*R + DIST*DIST - 0.5) / (2.0*R*DIST)
      IF(TEMP.GT.1.0) GOTO 20
      beta=2.0*ACOS(TEMP)

```

```

***** B.C. - Ray path starts at (I,0) and ends at (0,J) *****
                IF(K.EQ.I .AND. L.EQ.0)          beta=beta/2
                IF(K.EQ.0 .AND. L.EQ.J)          beta=beta/2

***** Chord (arc) length *****
                CLEN=R*beta
                IF(CLEN.LT.CMIN)          GOTO 20

***** Summation of parameter times *****
                TSUM(K,L)=TSUM(K,L)+TIME(I,J)*(CLEN/(C0+L*C1))/T est

***** Summation of chord (arc) lengths *****
                CSUM(K,L)=CSUM(K,L)+CLEN
20 CONTINUE
30 CONTINUE

***** Average slowness at all points *****
                DO 40 K=0,MROW
                DO 40 L=0,NCOL
                    IF(CSUM(K,L).EQ.0.0) THEN
                        SLOW(K,L)=0.0
                    ELSE
                        SLOW(K,L)=TSUM(K,L)/CSUM(K,L)
                    END IF
40 CONTINUE

***** Constrained parameters *****
                IF(LCONST) CALL PARCON(NCOL)

***** For each ray path... *****
                DO 70 I=1,M
                DO 70 J=1,N
                    IF(TIME(I,J).LE.0.0) GOTO 70
                    TCALC(I,J)=0.0

                RAYPAR=2.0*C1*I / SQRT((C0*C0+(C0+J*C1)**2+C1*C1*I*I) **2 - 4*C0
* *C0*(C0+C1*J)**2 )
                R=1.0/(RAYPAR*C1)
                alpha=2*ASIN(0.5*SQRT(FLOAT(I*I + J*J))/R)
                angle i= ATAN(I/FLOAT(J)) - alpha/2.0
                RLEN= R*alpha

```

```

NSEG= INT(RLEN/DELTAR)
SEG= RLEN/NSEG
Xo=-R*COS(angle i + alpha)
Zo=-R*SIN(angle i)

***** Interpolate slowness at midpoints of segment *****

DO 60 S=SEG/2.0,RLEN,SEG
      beta=S/R
      X=R*COS(ANGLE I + BETA) + XO
      Z=R*SIN(ANGLE I + BETA) + ZO
      TCALC(I,J)=TCALC(I,J) + SEG*CARINT(X,Z)
60 CONTINUE
70 CONTINUE

***** Mean square error *****

ERROR=0.0
DO 80 I=1,MROW
DO 80 J=1,NCOL
      IF(TIME(I,J).LE.0.0) GOTO 80
      ERR=TCALC(I,J)-TIME(I,J)
      ERROR=ERROR+ERR*ERR
80 CONTINUE
RETURN
END

SUBROUTINE ARTLB2 (MROW,NCOL,DELTAR,CMIN,C0,C1,ERROR)
*****
*
* ARTLB2: Algebraic Reconstruction Technique for a Linear
* velocity Background (ie. curved raypaths)
* Iterative portion
*
* Programmed by KLS--May 1984
*
* ***** CALLING VARIABLES *****
*
* MROW # of sources along x-axis
* NCOL # of receivers along z-axis
* DELTAR Approx. length of ray segment for integration
* CMIN Minimum acceptable chord length to use to calculate
* average slownesses
* C0 Velocity at surface
* C1 Change in velocity with a 'DX' increase in depth
*
*

```

```

*          ***** OUTPUT VARIABLES          *****
*
* ERROR   Mean square error of residual travel times
*
*          ***** COMMON STORAGE ***** INPUT *****
*
* LCONST  Logical switch for constraining the slowness parameters*
*          down the borehole (known from the sonic log)
*
* TIME()  First arrival times from the m th shot to the n th
*          geophone
*          Zero travel time signifies a lack of data
*
*          ***** COMMON STORAGE ***** I/O *****
*
* SLOW()  Slowness array
*
* TCALC() Calculated travel times

```

```

*****
PARAMETER      (M=20,N=10)
LOGICAL        LCONST
DIMENSION      CSUM(0:M,0:N),  ESUM(0:M,0:N)
COMMON/CONSTR/ LCONST
COMMON/VELOCI/ SLOW(0:M,0:N)
COMMON/TIMES/  TIME(M,N)
COMMON/CALCUL/ TCALC(0:M,0:N)

```

```

*****
*          ***** VARIABLE USAGE ***** VARIABLE USAGE *****
*
* Alpha   Angle of ray path
* Angle i Angle of incidence for ray at K,0
* Beta    Angle of intersection
* CSUM()  Sum of chord lengths around M,N parameter
* CLEN    Chord (arc) length intersecting parameter circle
* DIST    Distance between parameter point and origin of
*          the radius of curvature for the raypath
*
* ERR     Residual error
* ESUM()  Sum of errors associated with parameter M,N
* I       Position of source
* J       Position of reciever
* K       X parameter coordinate
* L       Z parameter coordinate
* NSEG    # of ray segments
* R       The radius of curvature for the ray
* RAYPAR  Ray parameter for the ray from iTH shot to jTH geophone*
* RLEN    Length of ray path from ith source to jth receiver

```

```

*      SEG      True length of ray path      *
*      X,Z      Midpoint of segment for integration *
*      Xo,Zo    Origin of the radius of curvature *
*              always a negative value      *
*
*      ***** UNITS OF DIMENSIONS *****
*
*      Angles      in radians
*      Distance    in grid units (spacing between geophones)
*      Time        in seconds
*      Slowness    in seconds/grid unit
*      Velocity    in grid units/second
*
*****
***** Clear summations *****

DO 10  K=0,MROW
DO 10  L=0,NCOL
      CSUM(K,L)=0.0
      ESUM(K,L)=0.0
10 CONTINUE

***** For each ray path... *****

DO 30  I=1,MROW
DO 30  J=1,NCOL

***** For no data on ray *****

      IF(TIME(I,J).LE.0.0) GOTO 30
      ERR=TIME(I,J)-TCALC(I,J)

***** Use geometry *****

RAYPAR=2.0*C1*I / SQRT((C0*C0+(C0+J*C1)**2+C1*C1*I*I) **2 - 4*C0
* *C0*(C0+C1*J)**2 )
      R=1.0/(RAYPAR*C1)
      alpha=2.0*ASIN(0.5*SQRT(FLOAT(I*I + J*J))/R)
      angle i=ATAN(I/FLOAT(J)) - alpha/2.0
      RLEN=R*alpha
      Xo=-R*COS(angle i + alpha)
      Zo=-R*SIN(angle i)

***** For each parameter circle... *****

DO 20  K=0,I
      L1=NINT(Zo + SQRT((R-1./SQRT(2.))**2 - (K-Xo)**2) + .5)

```

```

                IF(L1.LT.0)      L1=0
                L2=INT(Zo + SQRT((R+1./SQRT(2.))**2 - (K-Xo)**2))
                IF(L2.GT.J)      L2=J
DO 20  L=L1,L2

*      Find intersection of two equations
*      Equation one: Circle (X-K)**2 + (Z-L)**2 = 1/2
*      Equation two: Circle (X-Xo)**2 + (Z-Zo)**2 = R**2

***** Ray intersects parameter circle ??? *****

                DIST=SQRT((K-Xo)**2+(L-Zo)**2)
                IF(ABS(DIST - R) .GE. 1./SQRT(2.0))      GOTO 20

***** Locate intersection angle *****

***** Law of cosines *****

***** TEMP avoids computer roundoff errors *****

                TEMP=(R*R + DIST*DIST - 0.5) / (2.0*R*DIST)
                IF(TEMP.GT.1.0) GOTO 20
                beta=2.0*ACOS(TEMP)

***** B.C. - Ray path starts at (I,0) and ends at (0,J) *****

                IF(K.EQ.I .AND. L.EQ.0)      beta=beta/2
                IF(K.EQ.0 .AND. L.EQ.J)      beta=beta/2

***** Chord (arc) length *****

                CLEN=R*beta
                IF(CLEN.LT.CMIN)      GOTO 20

***** Summation of chord (arc) lengths *****

                CSUM(K,L)=CSUM(K,L) + CLEN

***** Summation of error around parameter *****

                ESUM(K,L)=ESUM(K,L) + ERR*CLEN/RLEN

20      CONTINUE
30      CONTINUE

***** Update parameter matrix *****

DO 40  K=0,MROW

```

```

DO 40  L=0,NCOL
      IF(CSUM(K,L).EQ.0.0) THEN
          SLOW(K,L)=0.0
      ELSE
          SLOW(K,L)=SLOW(K,L) + ESUM(K,L)/CSUM(K,L)
      END IF
40    CONTINUE

*****  Constrained parameters  *****

      IF(LCONST) CALL PARCON(NCOL)

*****  For each ray path...  *****

      DO 70  I=1,M
      DO 70  J=1,N
      IF(TIME(I,J).LE.0.0) GOTO 70
      TCALC(I,J)=0.0

      RAYPAR=2.0*C1*I / SQRT((C0*C0+(C0+J*C1)**2+C1*C1*I*I) **2 - 4*C0
* *C0*(C0+C1*J)**2 )
      R=1.0 / (RAYPAR*C1)
      alpha= 2*ASIN(0.5*SQRT(FLOAT(I*I + J*J))/R)
      angle i=ATAN(I/FLOAT(J)) - alpha/2.0
      RLEN= R*alpha
      NSEG= INT(RLEN/DELTAR)
      SEG= RLEN/NSEG
      Xo=-R*COS(angle i + alpha)
      Zo=-R*SIN(angle i)

*****  Interpolate slowness at midpoints of segment  *****

      DO 60  S=SEG/2.0,RLEN,SEG
          beta=S/R
          X=R*COS(ANGLE I + BETA) + XO
          Z=R*SIN(ANGLE I + BETA) + ZO
          TCALC(I,J)=TCALC(I,J) + SEG*CARINT(X,Z)
60    CONTINUE
70    CONTINUE

*****  Mean square error  *****

      ERROR=0.0
      DO 80  I=1,MROW
      DO 80  J=1,NCOL
          IF(TIME(I,J).LE.0.0) GOTO 80
          ERR=TCALC(I,J)-TIME(I,J)
          ERROR=ERROR+ERR*ERR

```

```

80      CONTINUE
        RETURN
        END

```

```

FUNCTION CARINT(X,Y)

```

```

*****
*
*      CARINT: CARtesian INTerpolation
*      Interpolates the value of F(X,Y)
*      Assumes function values are known at integer values of
*      X and Y — Which are stored in array U()
*
*      ***** CALLING VARIABLES *****
*
*      X,Y      Coordinates of point to interpolate
*
*      ***** COMMON STORAGE ***** INPUT *****
*
*      U()      Values of function at integer values for X,Y
*
*****
PARAMETER      (M=20,N=10)
COMMON/VELOC/  U(0:M,0:N)
I=INT(X)μ      J=INT(Y)
P=X-FLOAT(I)μ  Q=Y-FLOAT(J)
F00=U(I,J)μ    F10=U(I+1,J)
F01=U(I,J+1)μ  F11=U(I+1,J+1)

***** Two point formula *****

IF(.NOT.(F11.EQ.0.0.AND.F10.EQ.0.0)) GOTO 10
CARINT=(1-Q)*F00 + Q*F01
RETURN

***** Three point formula *****

10      IF(F11.NE.0.0) GOTO 20
CARINT=(1.-P-Q)*F00 + P*F10 + Q*F01
RETURN

***** Three point formula *****

20      IF(F00.NE.0.0) GOTO 30
CARINT=(1-P)*F01 + (1-Q)*F10 + (P+Q-1)*F11
RETURN

```

```

***** Four point formula *****

30 CARINT=(1.-P)*(1.-Q)*F00+P*(1.-Q)*F10+Q*(1.-P)*F01+P*Q*F11
RETURN
END

SUBROUTINE PARCON(NCOL)
IMPLICIT LOGICAL(L)
PARAMETER (M=20,N=10)
COMMON/CONSTR/ LCONST,VELVAR,CONST(N)
COMMON/VELOCI/ SLOW(0:M,0:N)
*****
*
* PARCON: PARAMeter CONstraining routine
* Fixes the slowness parameter values down the borehole
* as determined by the sonic log
*
* ***** CALLING VARIABLES *****
*
* NCOL # of geophones positions down the well
*
* ***** COMMON STORAGE ***** INPUT *****
*
* CONST() Constrained parameter values
*
* LCONST Logical switch
*
* VELVAR Allowable velocity variation from the fixed value
* Deviation in percentage from fixed value
*
* ***** COMMON STORAGE ***** I/O *****
*
* SLOW() Slowness matrix
*
*****
DO 10 J=1,NCOL
IF(CONST(J).LE.0.0) GOTO 10
VARY=CONST(J)*VELVAR/100.
IF(SLOW(0,J).LT.CONST(J)-VARY) SLOW(0,J)=CONST(J)-VARY
IF(SLOW(0,J).GT.CONST(J)+VARY) SLOW(0,J)=CONST(J)+VARY
10 CONTINUE
RETURN
END

FUNCTION VEL (DX,SLOW)
*****

```

```

*      VEL:      Computes velocity from the slowness      *
*****
      IF(SLOW.LE.0.0) THEN
          VEL=0.0
      ELSE
          VEL=DX/SLOW
      END IF
      RETURN
      END

      SUBROUTINE ABBREV      (CMMND,NUMCOM,WORD,NUMLET)
      IMPLICIT      LOGICAL(L)
      CHARACTER      WORD*5, CMMND*65
*****
*
*      ABBREV: Expands abbrev. commands to their full length
*
*      *****      CALLING VARIABLES      *****
*
*      CMMND      Legal command strings
*      NUMCOM      # of legal commands
*      WORD      Inputed command string
*      NUMLET      # of letters in inputed command
*
*      *****      OUTPUT VARIABLES      *****
*
*      WORD      The abbrev command expanded to its full length
*
*****
      CALL UCASE2      (.TRUE.,NUMLET,WORD)
      DO 20      I=NUMLET+1,NUMLET*NUMCOM+NUMLET,NUMLET
          LBEFOR=.FALSE.
          LAFTER=.FALSE.
      DO 10      J=1,NUMLET
          IF(WORD(J:J).EQ.' ')      GOTO 30
          K=I+J-1
          IF(WORD(J:J).NE.CMMND(K:K))      GOTO 20
          IF(WORD(J:J).NE.CMMND(K-NUMLET:K-NUMLET)) LBEFOR=.TRUE.
          IF(WORD(J:J).NE.CMMND(K+NUMLET:K+NUMLET)) LAFTER=.TRUE.
10      CONTINUE
          RETURN
20      CONTINUE
          RETURN
30      IF(LBEFOR.AND.LAFTER)      WORD(1:NUMLET)=CMMND(I:I+NUMLET-1)
          RETURN
      END

```

```

SUBROUTINE UCASE2      (LCASE,NUMLET,WORD)
IMPLICIT LOGICAL(L)
CHARACTER WORD*5,UCALPH*26,LCALPH*26
*****
*
* UCASE2: Changes an alphanumeric string to a single case
*
* ***** CALLING VARIABLES *****
*
* LCASE .TRUE. Returns all upper case alphanumerics
*       .FALSE. Returns all lower case alphanumerics
* NUMLET # of letters in string
* WORD Alphanumeric string
*
* ***** OUTPUT VARIABLES *****
*
* WORD Alphanumeric string converted to upper/lower case
*
*****
UCALPH='ABCDEFGHIJKLMNOPQRSTUVWXYZ'
LCALPH='abcdefghijklmnopqrstuvwxyz'
DO 30 I=1,NUMLET
DO 20 J=1,26
IF(LCASE) THEN
IF(WORD(I:I).NE.LCALPH(J:J)) GOTO 10
WORD(I:I)=UCALPH(J:J)
GOTO 30
ELSE
IF(WORD(I:I).NE.UCALPH(J:J)) GOTO 10
WORD(I:I)=LCALPH(J:J)
GOTO 30
10 END IF
20 CONTINUE
30 CONTINUE
RETURN
END

```